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An Investigation of the 8 June 2007 Heavy-Rainfall Event in Central Taiwan

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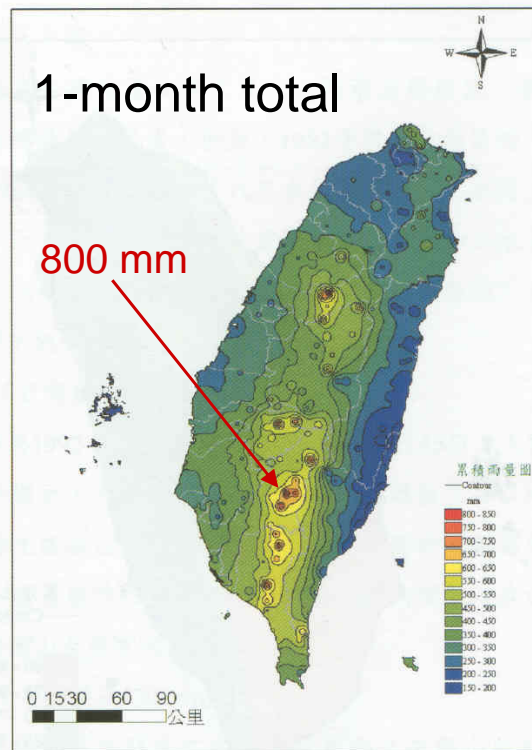
Presentation Outline:

- **Introduction**
 - **Case Review and Motivation**
 - **Synoptic Condition**
 - **Model and Experiment Design**
 - **Simulation Result**
 - **Discussion and Conclusion**
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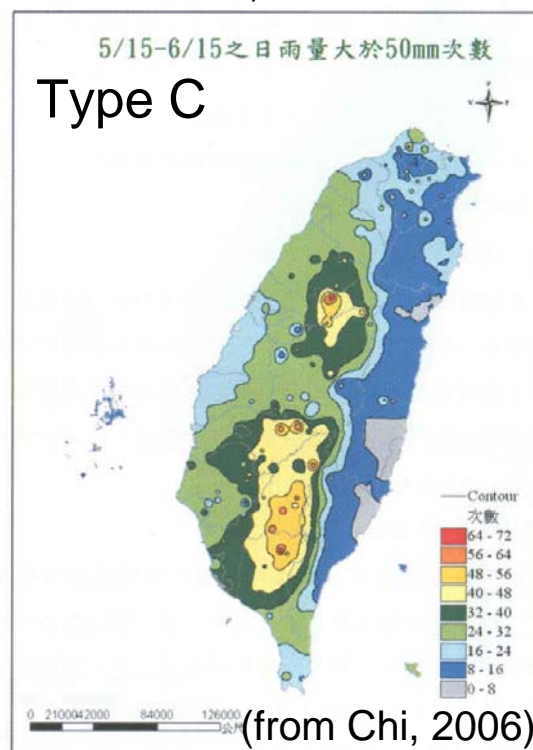
1. Introduction: (1)

- Heavy rainfall during the Mei-yu season in Taiwan occurs both over mountains and coastal regions

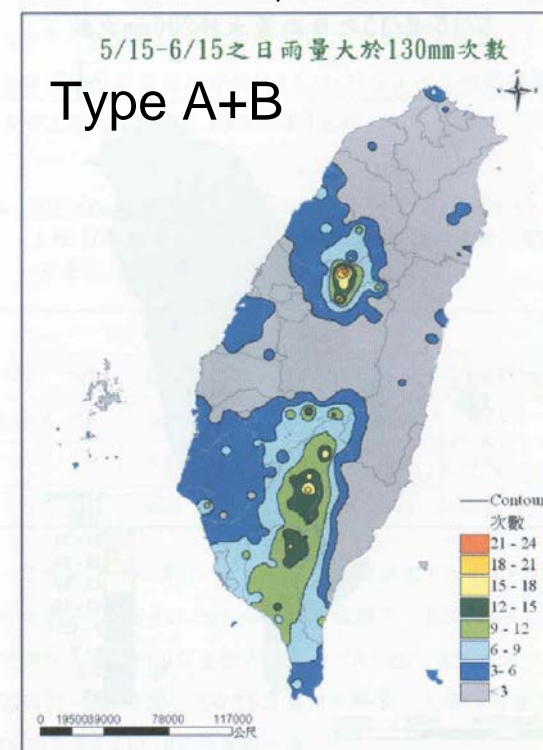
Mean total rainfall (mm)
ARMTS, 5/15-6/15, 1992-04



Total freq. of ≥ 50 mm/day
5/16-6/15, 1992-2004



Total freq. of ≥ 130 mm/day
5/16-6/15, 1992-2004

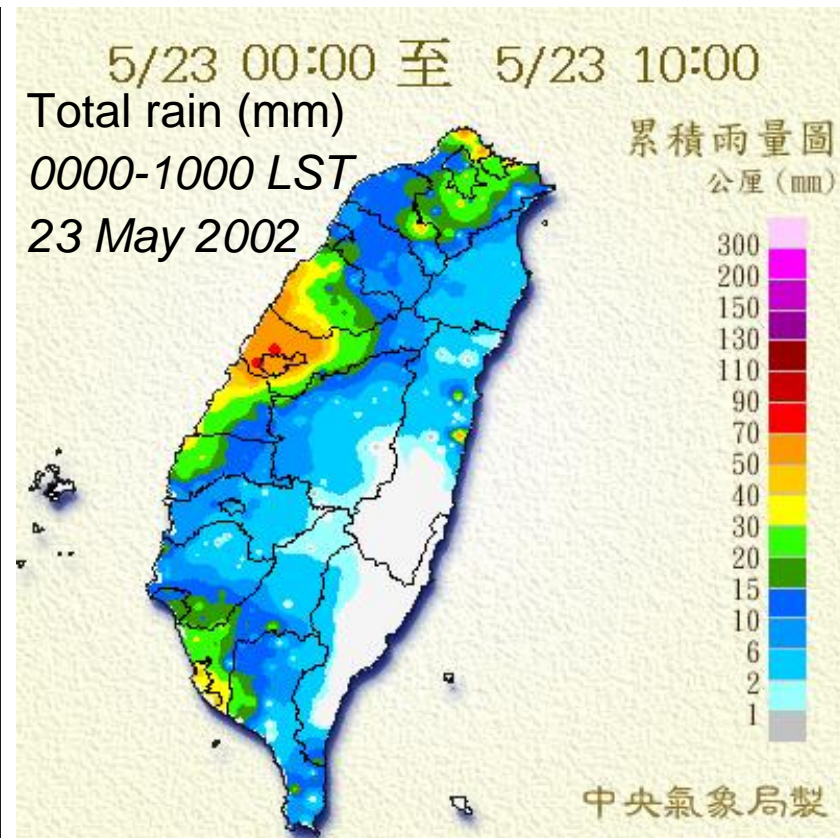
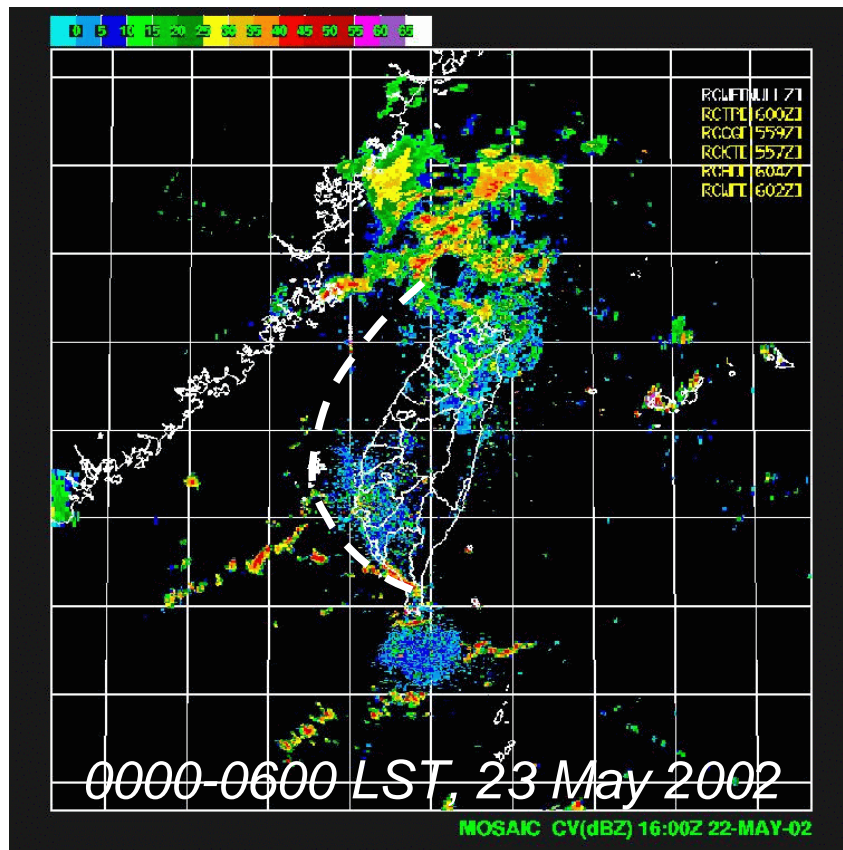


1. Introduction: (2)

- To accurately forecast the initiation, development, and dissipation of mesoscale convective systems (MCSs) responsible for heavy rainfall at correct time and location around Taiwan is often difficult
- Steep terrain and typically moist and unstable atmosphere
- MCSs and deep convection initiated by a variety of mechanisms: Frontal uplift, topographic uplift, dynamical forcing aloft, terrain blocking effect, solar heating during daytime (free convection), island circulation (land-sea and/or mountain-valley breezes), outflow boundary from previous/existing convection, ... etc.

1. Introduction: (3)

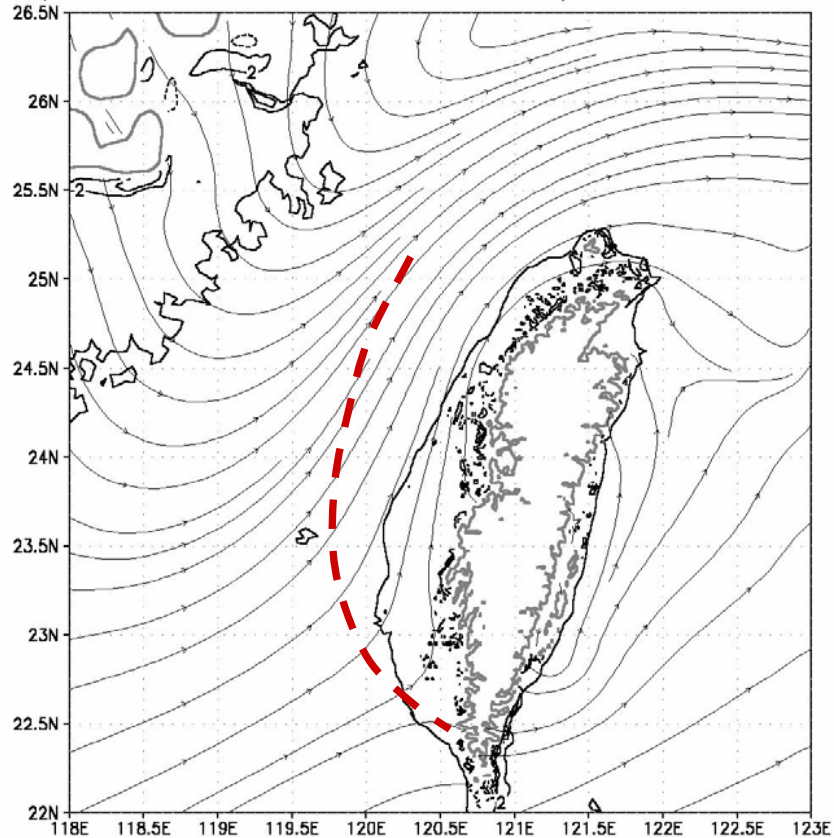
- An example of rainfall near the coast of central Taiwan, linked to terrain blocking of prevailing SW flow offshore



1. Introduction: (4)

- CReSS model simulation (Wang et al., 2005, MWR)

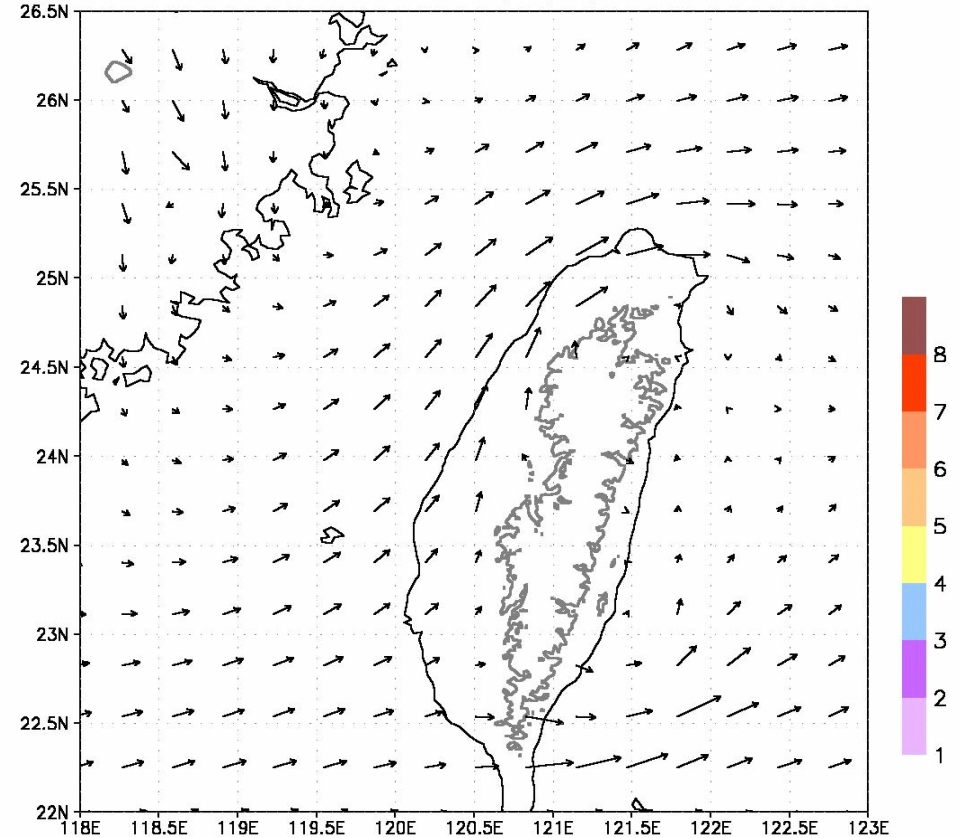
5/22 2300 LST FTRN strm/conv z=687m



Units: 10^{-4} s^{-1}

GRADS: COLA/IGES

5/23 0000 LST FTRN u/v/q2max z=1006m

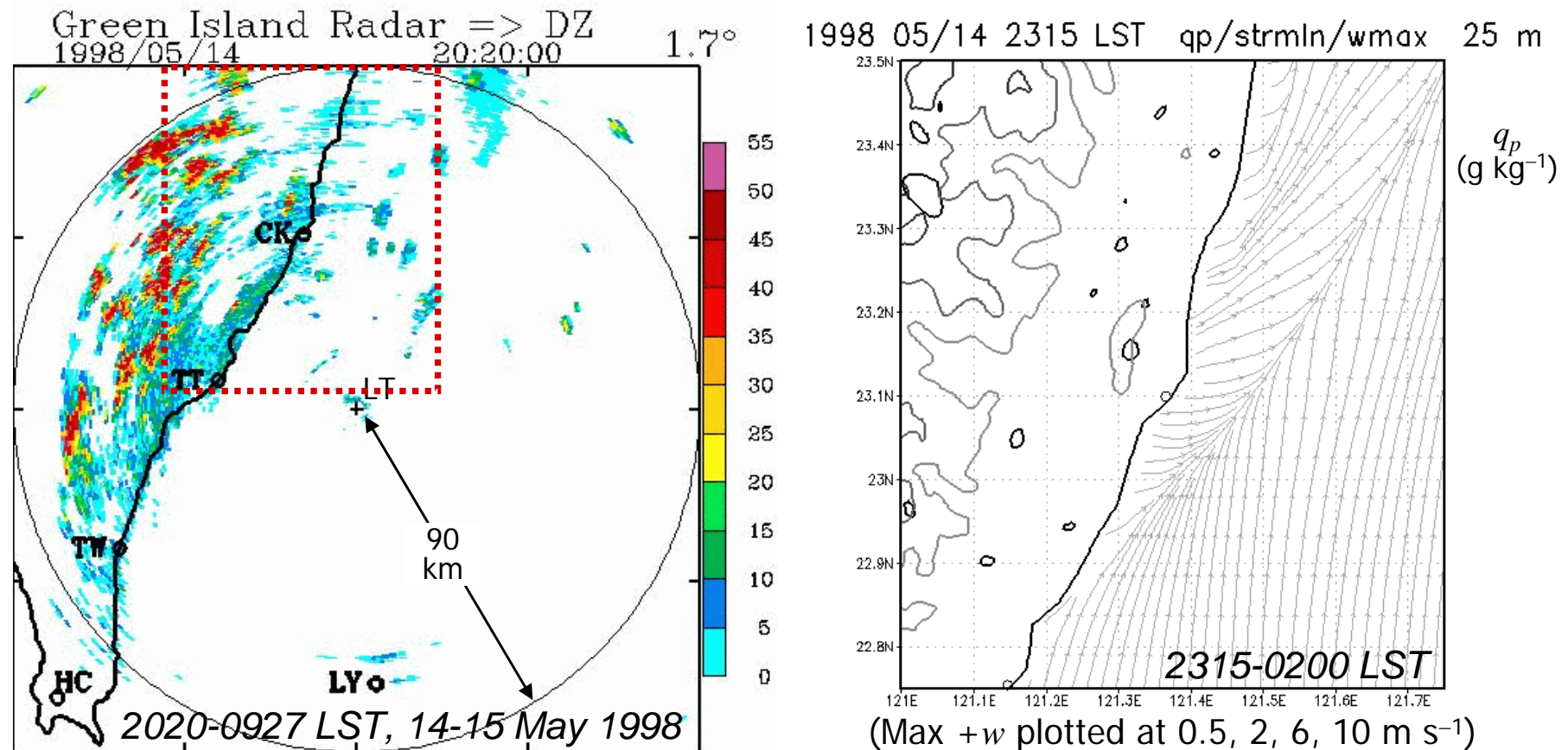


Units: g kg^{-1}

→ 20

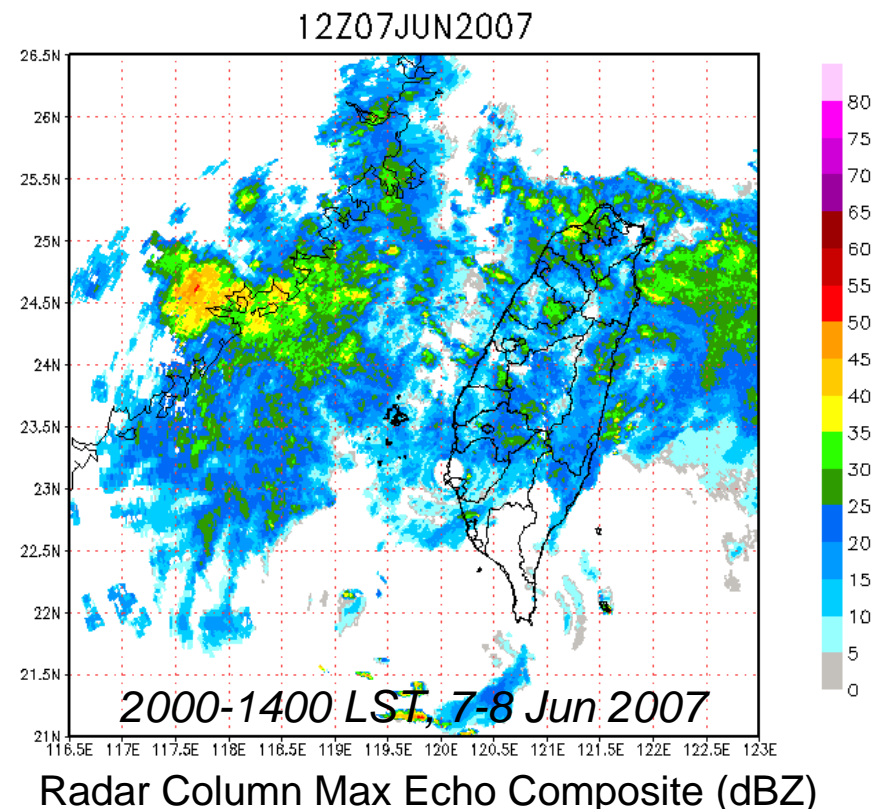
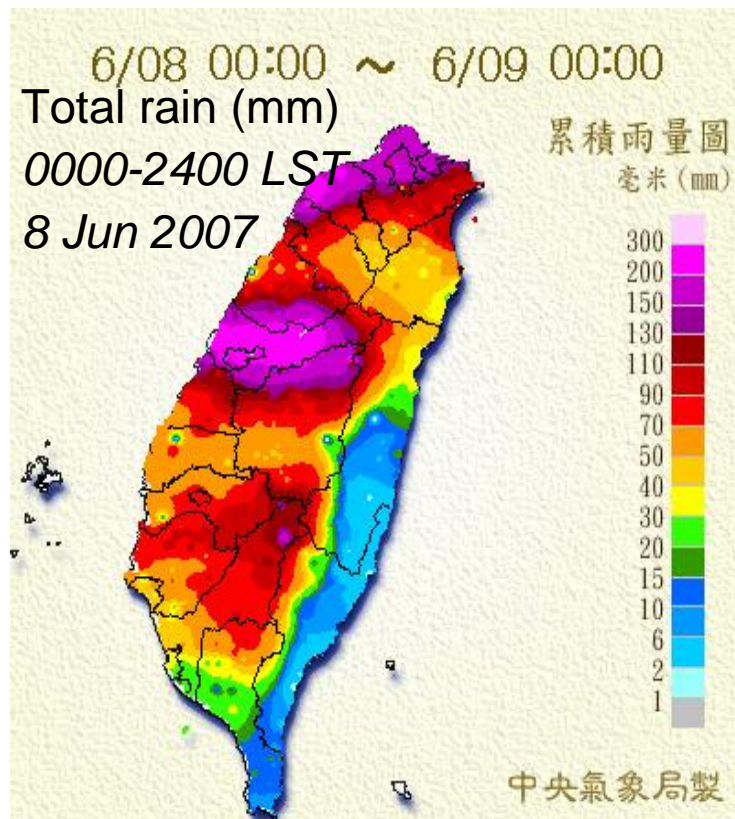
1. Introduction: (5)

- Another example of repeated convection initiation due to outflow boundary (Wang and Huang, 2009, GRL)

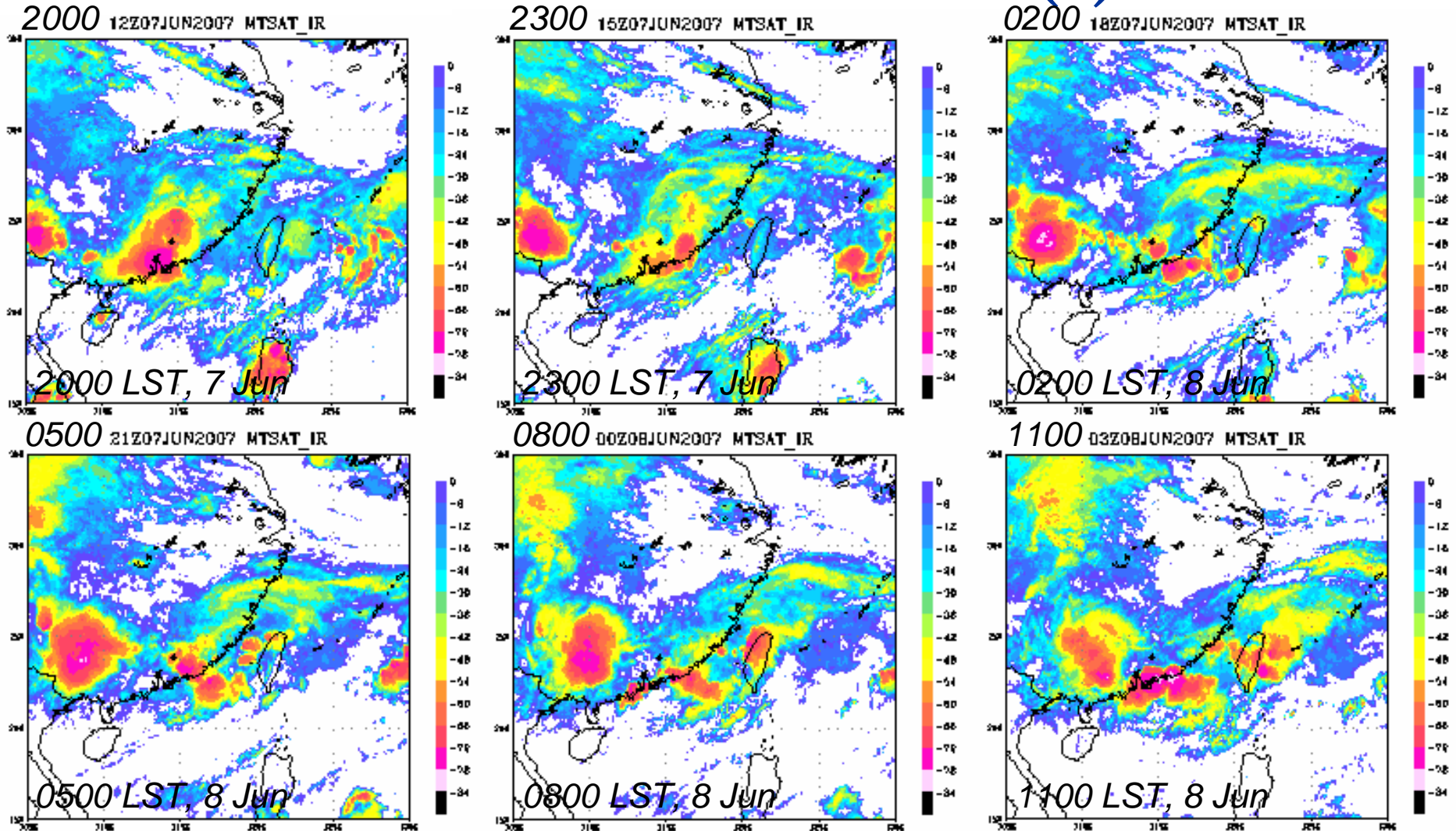


2. Case Review and Motivation: (1)

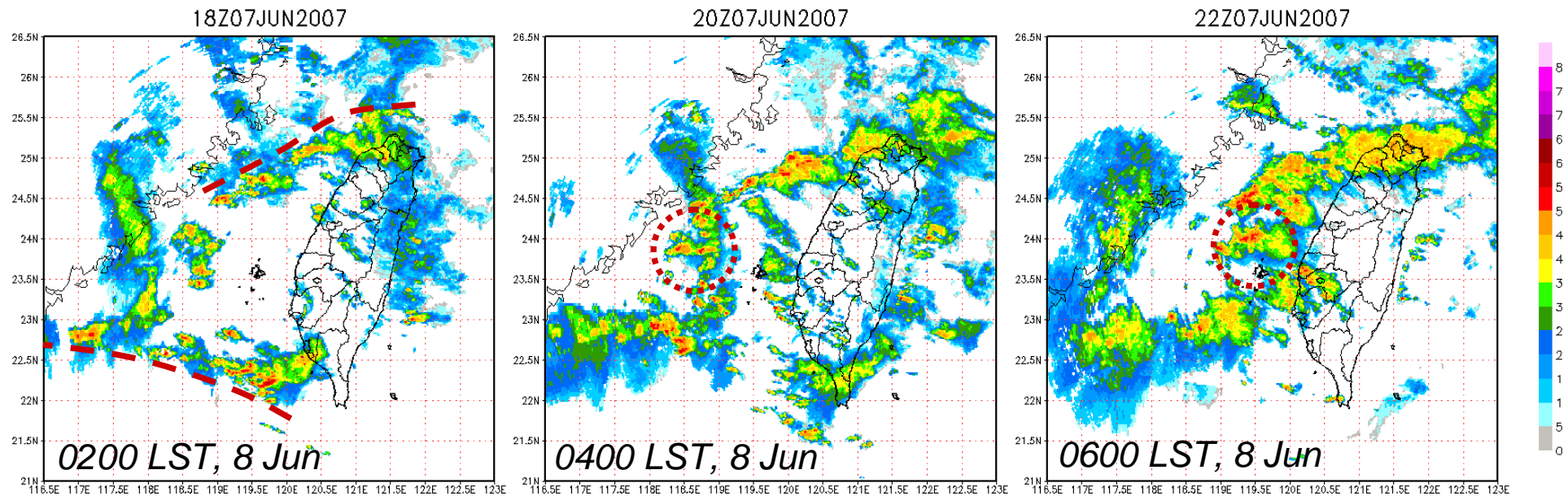
- Significant rainfall over northern, central, and southern Taiwan on 8 Jun 2007, but heaviest in central Taiwan, including the coastal region (daily total >250 mm)



2. Case Review and Motivation: (2)



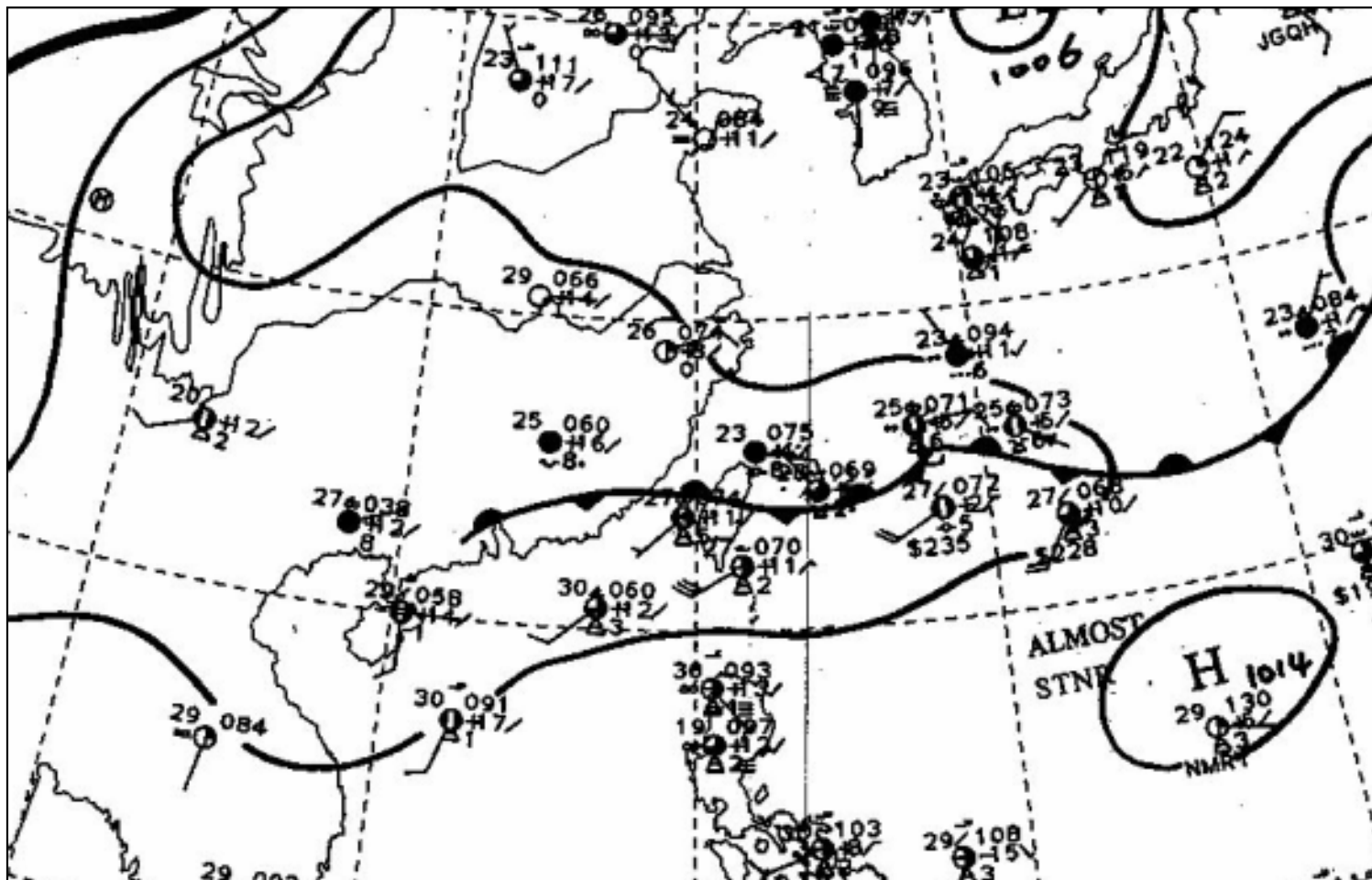
2. Case Review and Motivation: (3)



- Two pre-existing lines of convection: One over northern strait along the front (ENE-WSW) and the 2nd over southern strait (ESE-WNW) linking to MCSs moving offshore from China
- Why was convection initiated in between the two lines (over central Taiwan Strait), to eventually move over central Taiwan and cause heavy rainfall along the coast?

3. Synoptic Condition: (1)

CWB Surface map: 0800 LST (0000 UTC), 8 Jun 2007

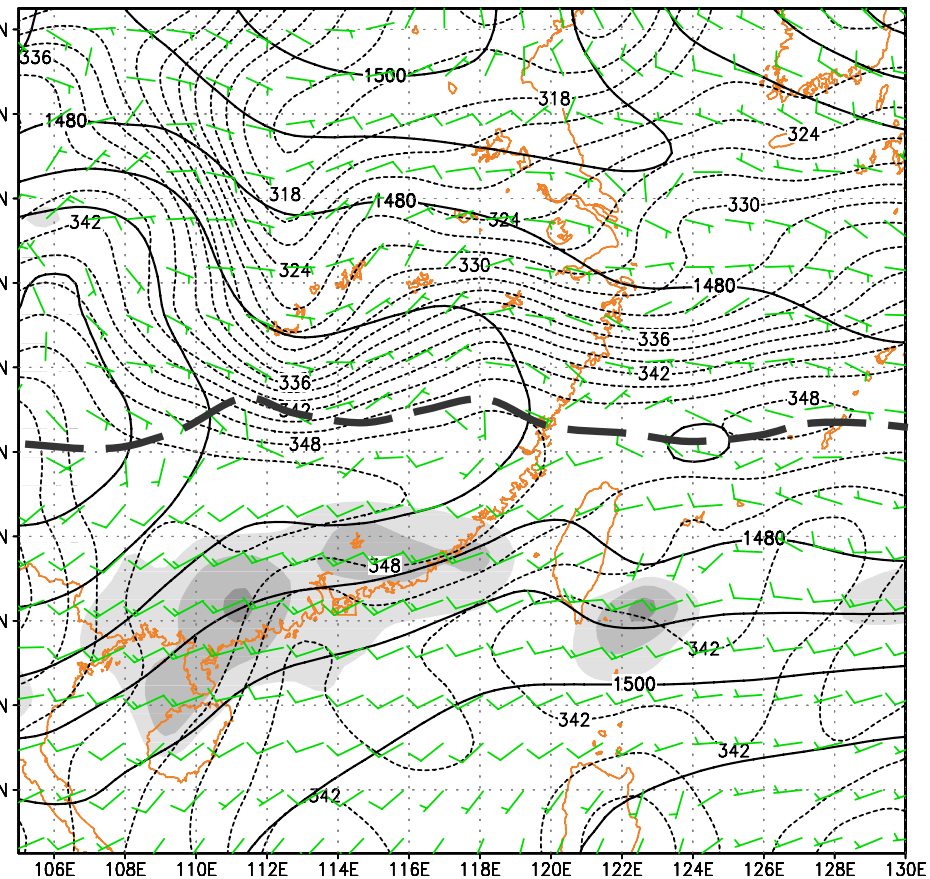
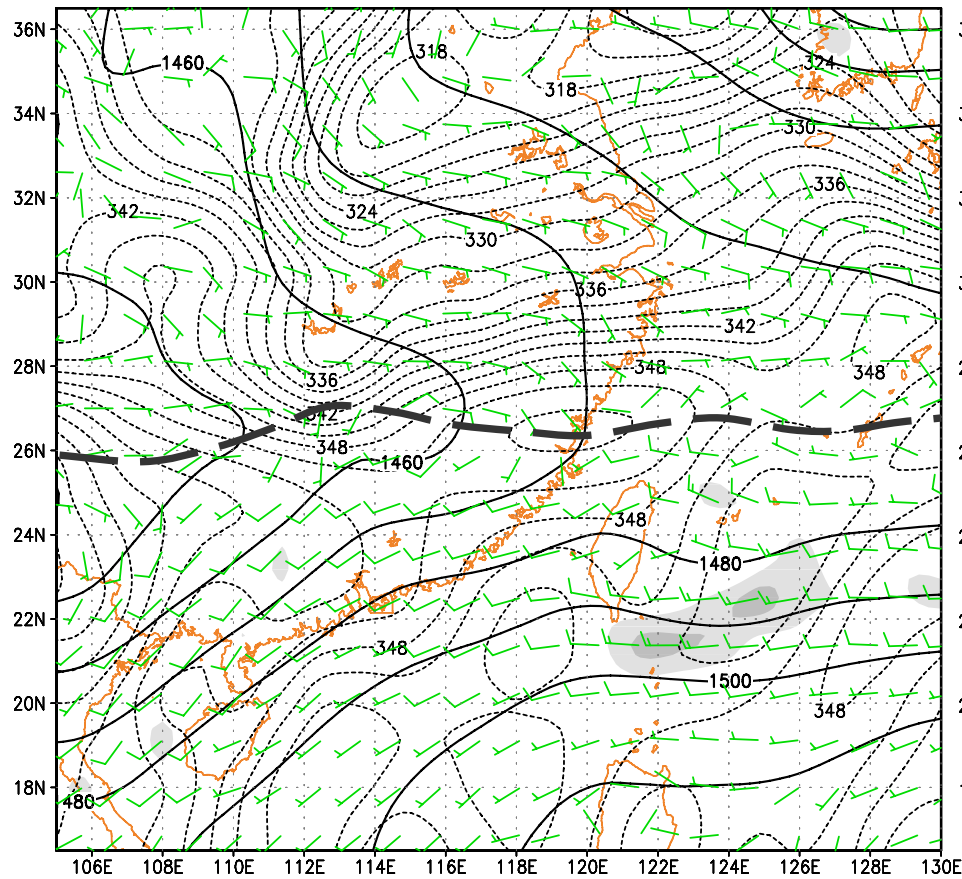


3. Synoptic Condition: (2)

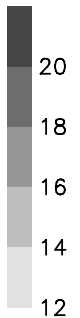
2000 LST, 7 Jun 2007 (850 hPa, ECMWF) 0800 LST, 8 Jun 2007

850hPa EPT&Wind&H 12Z07JUN2007

850hPa EPT&Wind&H 00Z08JUN2007



Wind speed (m/s)

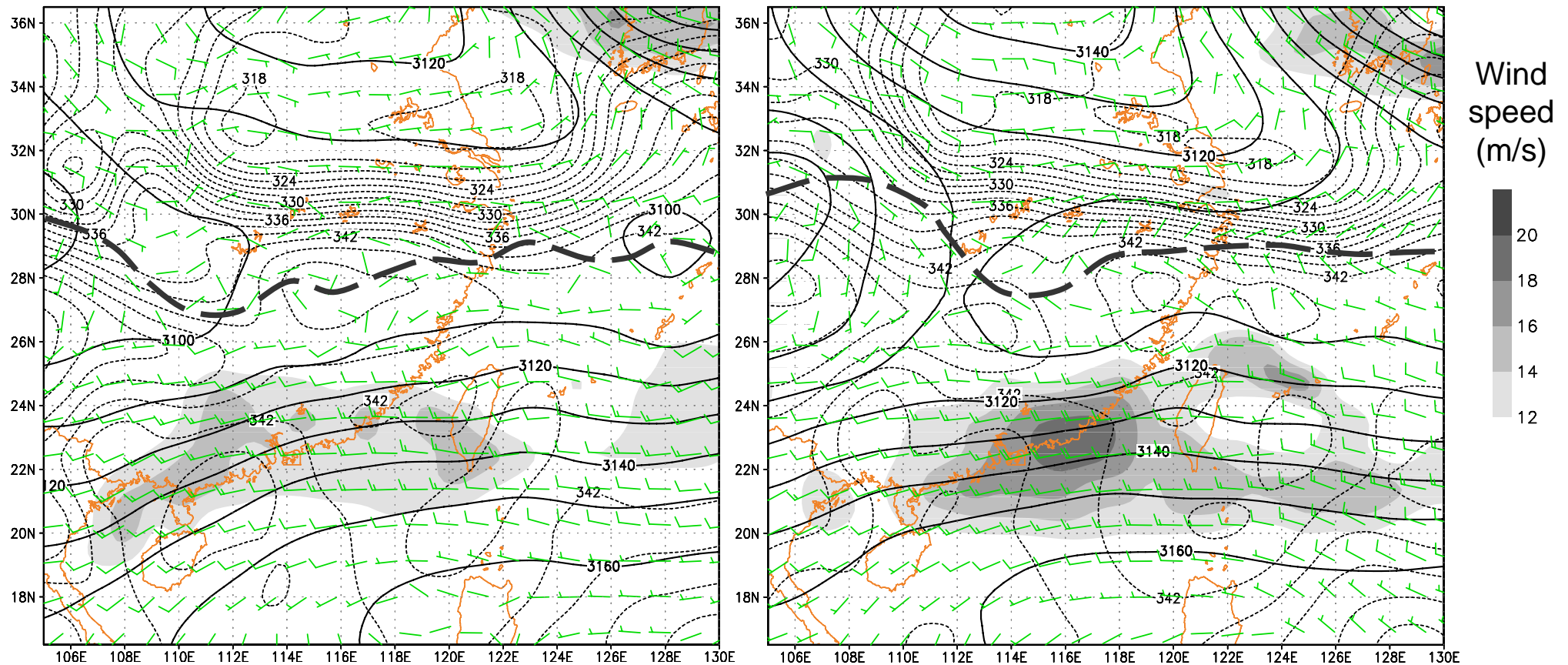


3. Synoptic Condition: (3)

2000 LST, 7 Jun 2007 (700 hPa, ECMWF) 0800 LST, 8 Jun 2007

700hPa EPT&Wind&H 12Z07JUN2007

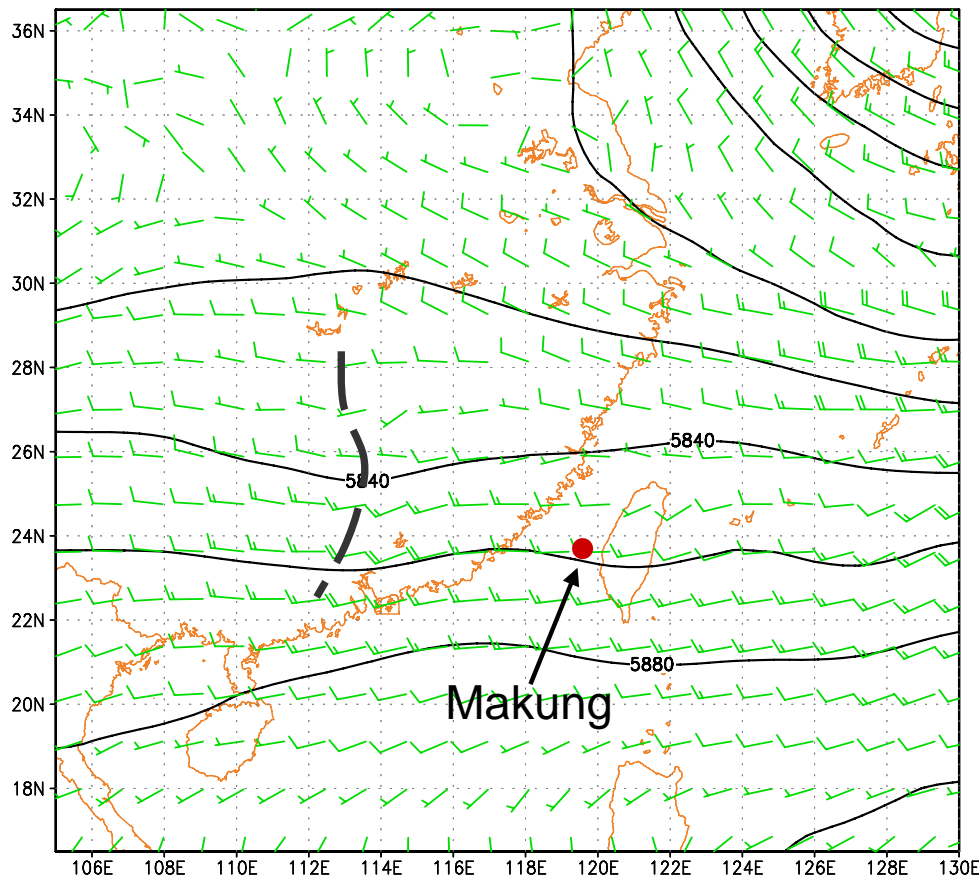
700hPa EPT&Wind&H 00Z08JUN2007



3. Synoptic Condition: (4)

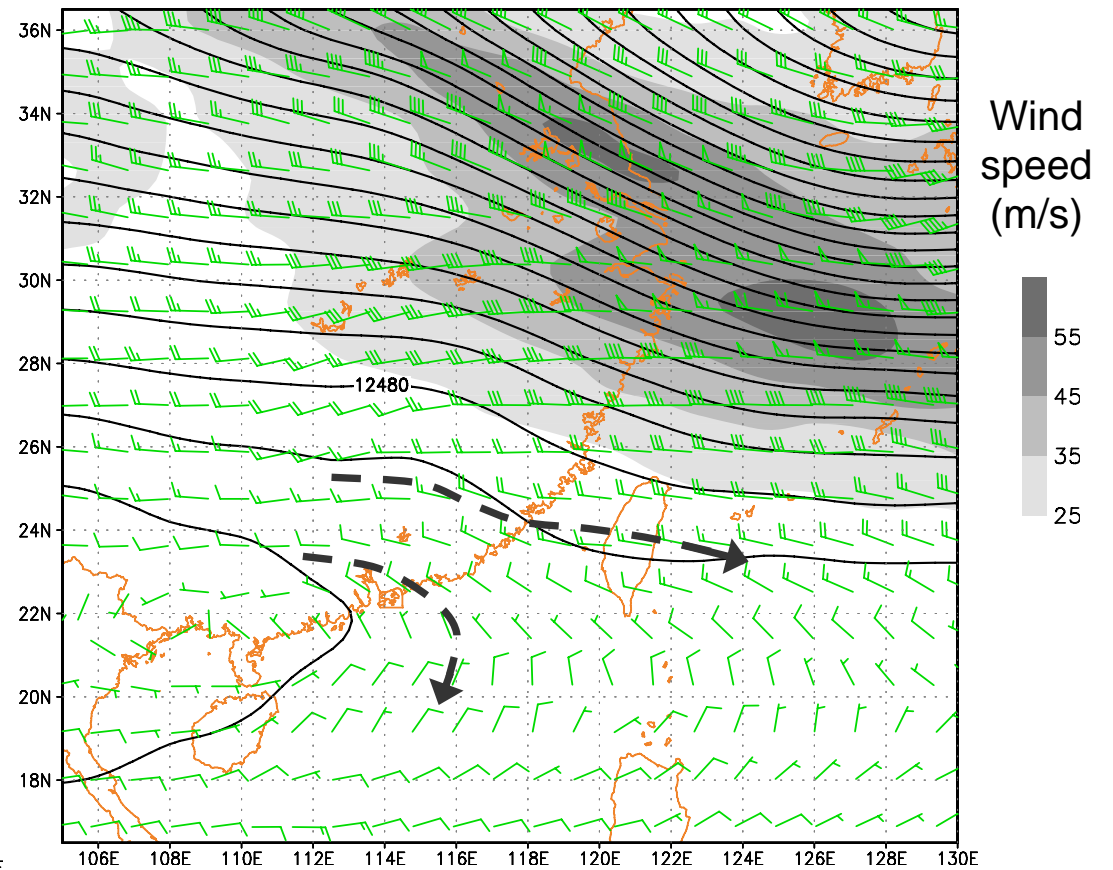
500 hPa, 2000 LST, 7 Jun 2007

500hPa H&Wind 12Z07JUN2007



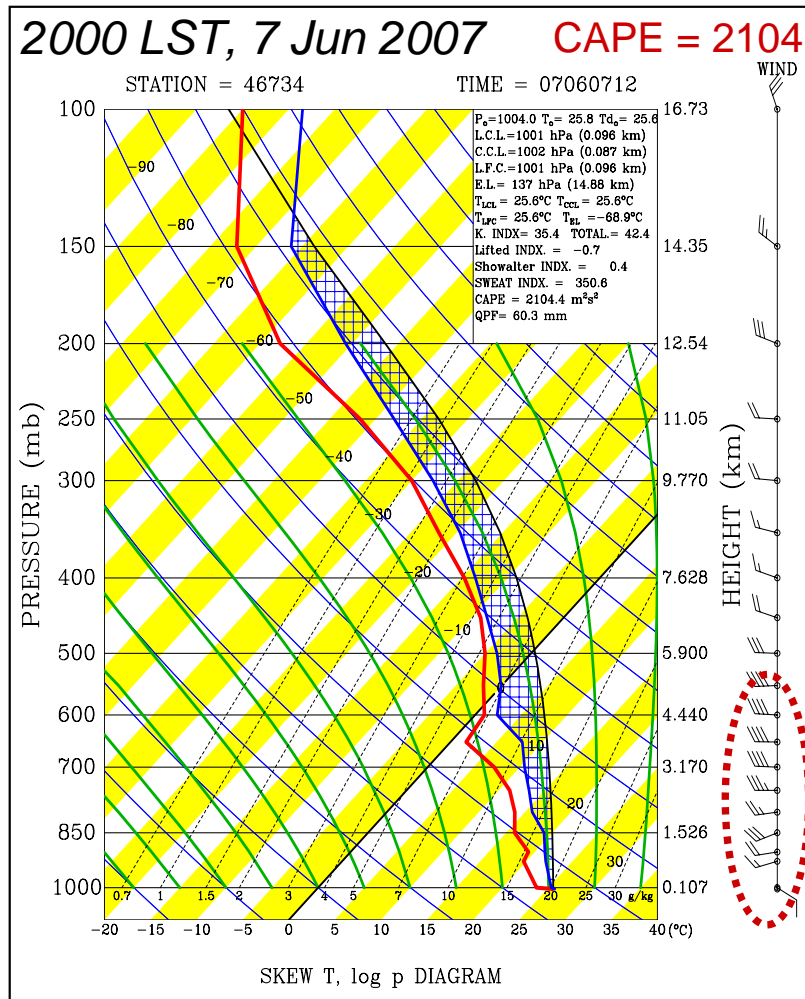
200 hPa, 2000 LST, 7 Jun 2007

200hPa H&Wind 12Z07JUN2007

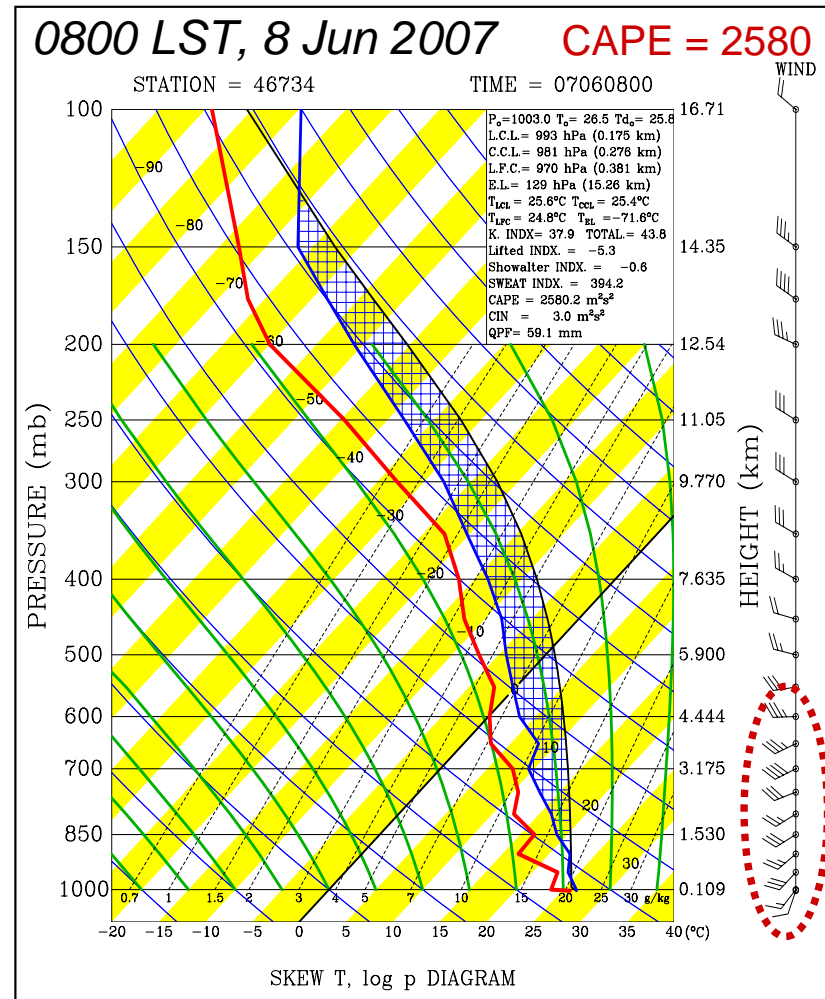


3. Synoptic Condition: (5)

Sounding at Makung (46734)



Sounding at Makung (46734)

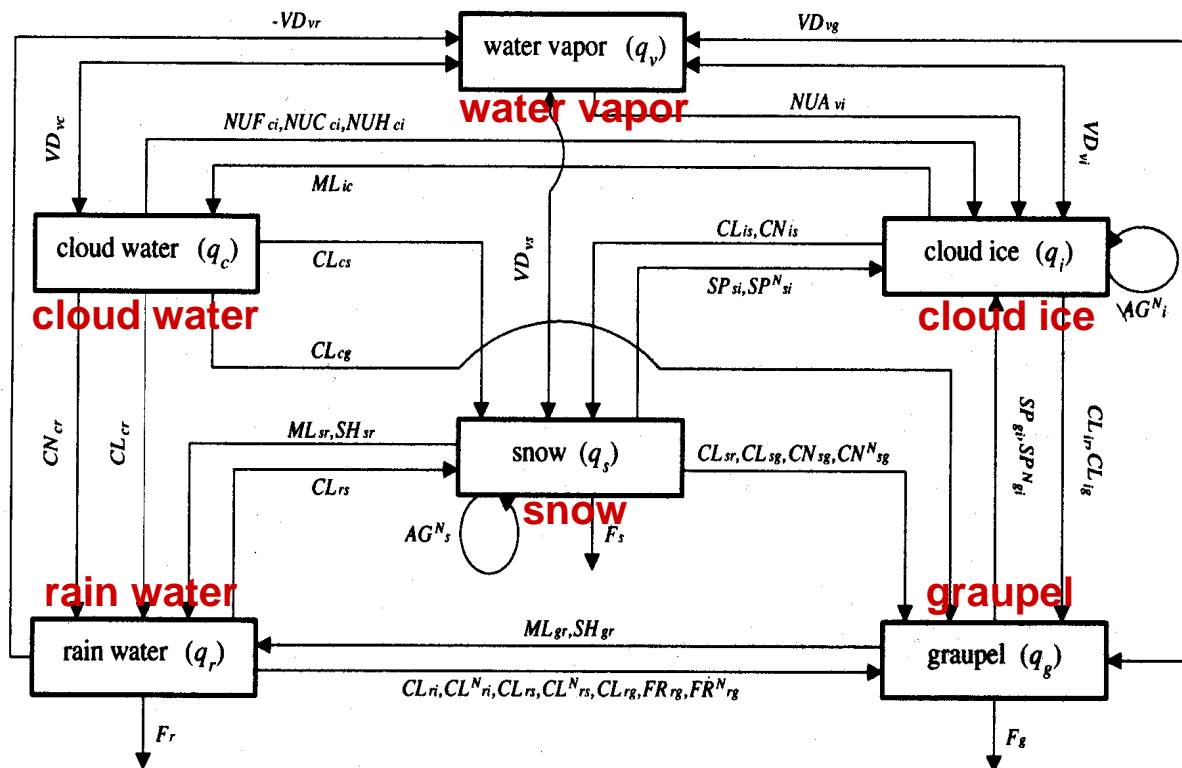


3. Synoptic Condition: (6)

- Synoptic-scale Mei-yu front running East-West and stationary over central Taiwan and central Taiwan Strait; front (shear-line) at 850 and 700 hPa farther north by about 300-400 km
- West-southwesterly Low-level jet (LLJ) of about 15 m s^{-1} at 850/700 hPa, providing strong vertical wind shear
- Short-wave trough approaching Taiwan at 500 hPa
- Diffluent northwesterly flow at 200 hPa over southern Taiwan Strait
- Central/southern Taiwan Strait ahead of the front, with high θ_e values, large convective available potential energy (CAPE, $>2000 \text{ m}^2 \text{ s}^{-2}$) and nearly no convective inhibition (CIN)
- Level of free convection (LFC) very close to surface, within 350 m (1001 and 970 hPa in Makung soundings)

4. Model and Experiment Design: (1)

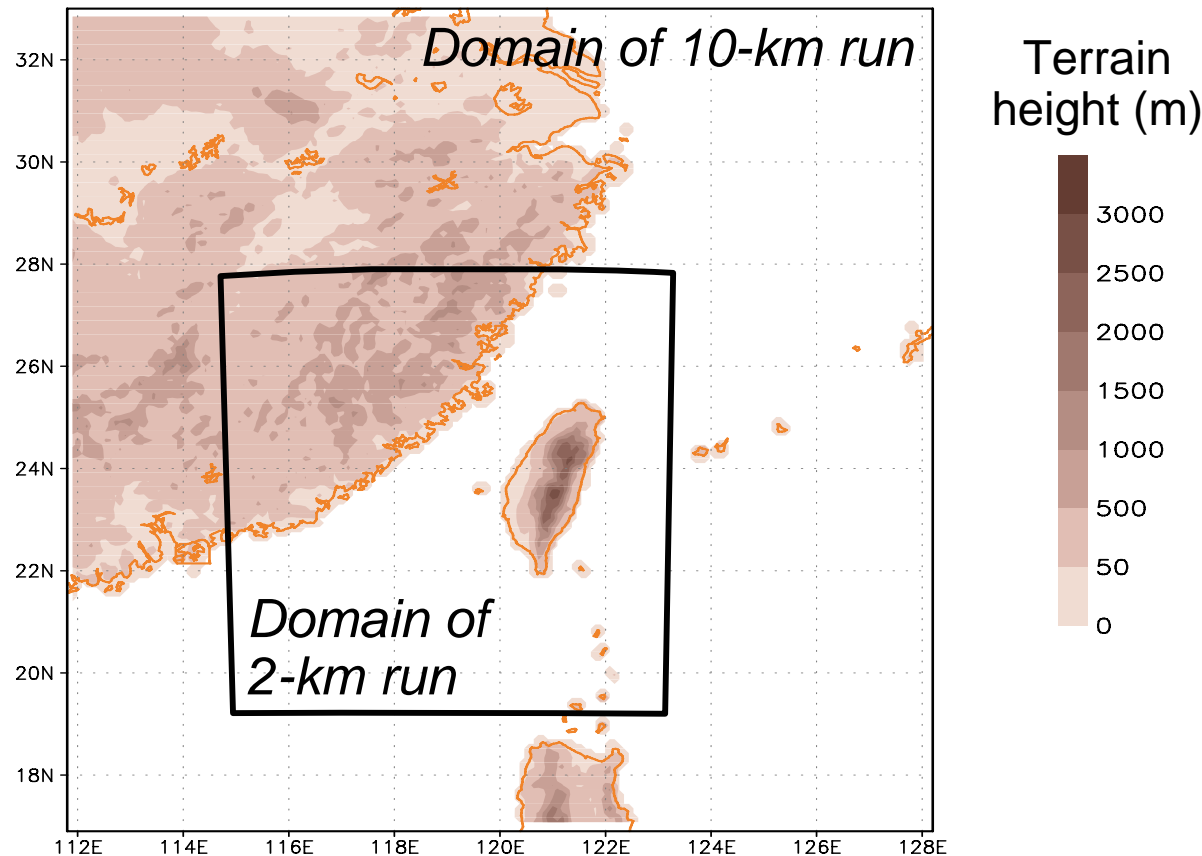
- Cloud-Resolving Storm Simulator (CReSS) of Nagoya Univ.
 - ❑ Non-hydrostatic, compressible, terrain following coordinate
 - ❑ Explicit cloud microphysics with no cumulus parameterization



- ❑ Nucleation:
 - Deposition/sorption, condensation-freezing, homogeneous process, contact, secondary
- ❑ Deposition, evaporation, sublimation, melting, freezing
- ❑ Falling, collection, conversion, aggregation, water shedding

4. Model and Experiment Design: (2)

- Two experiments are performed at increasingly higher resolution: 10-km and 2-km runs



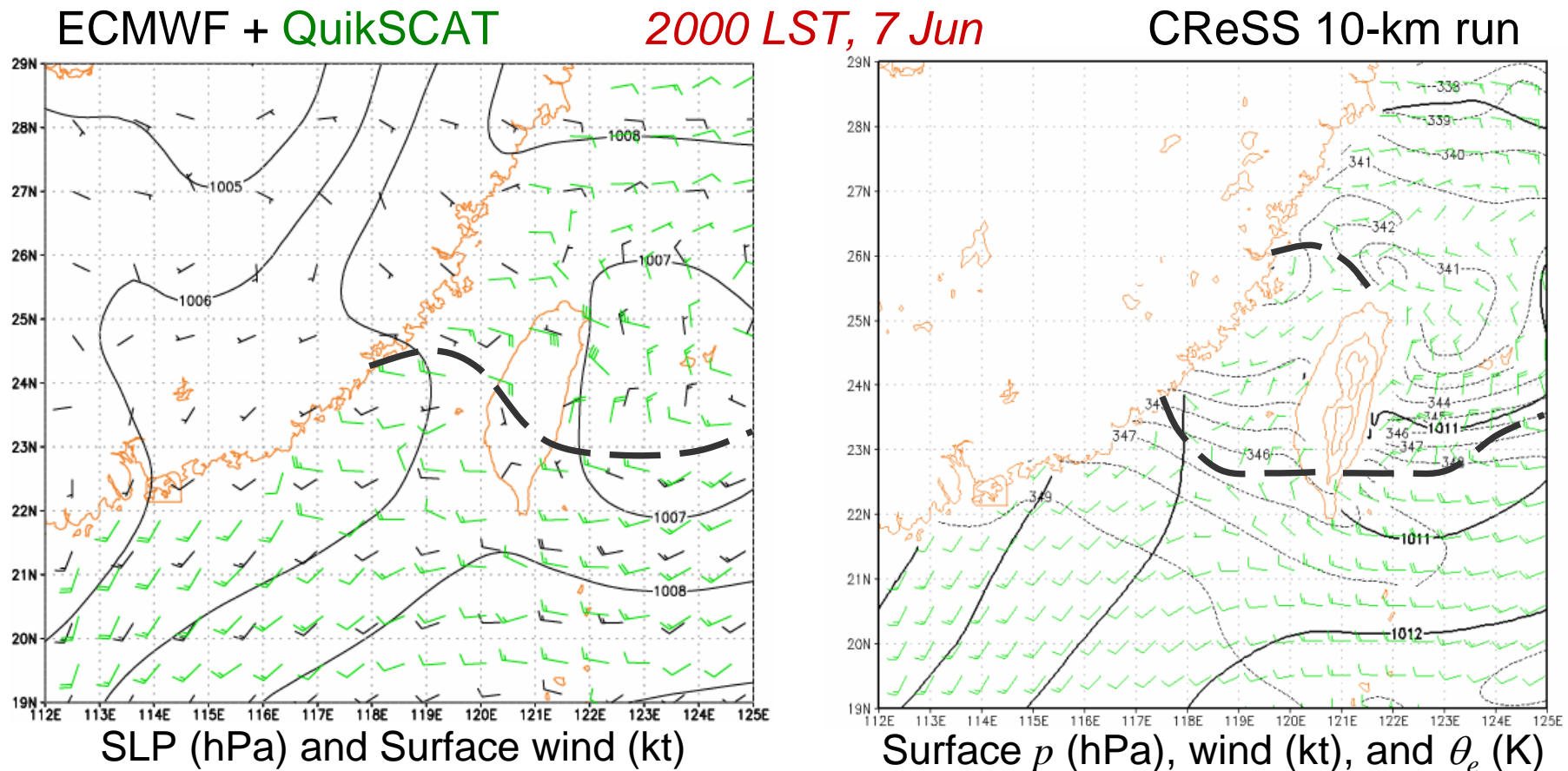
4. Model and Experiment Design: (3)

Experiment	One (10 km)	Two (2 km)
Projection	Lambert Conformal, center at 120°E, secant at 20°N and 50°N	
Grid spacing (km)	10 × 10 × 0.6*	2 × 2 × 0.35*
Dimension (x, y, z)	180 × 180 × 40	540 × 480 × 60
Minimum Δz	50 m	50 m (8 levels below 1 km)
Topography, SST	Real at (1/120)°, fixed at 28°C	
IC/LBCs	ECMWF 1.125°, 6 h	10-km run, 1 h
Initial time	2000 LST, 6 Jun	2000 LST, 6 Jun
Integration length and output frequency	72 h, 1 h	60 h, 10 min
Cloud microphysics	Bulk cold rain	
PBL parameterization	1.5-order closure with TKE prediction	
Surface processes	Energy/momentum fluxes, shortwave/longwave radiation	
Soil model	41 levels, every 5 cm to 2 m deep	

* The vertical grid spacing (Δz) of CReSS is stretched (smallest at the bottom), and the value given here is the averaged spacing.

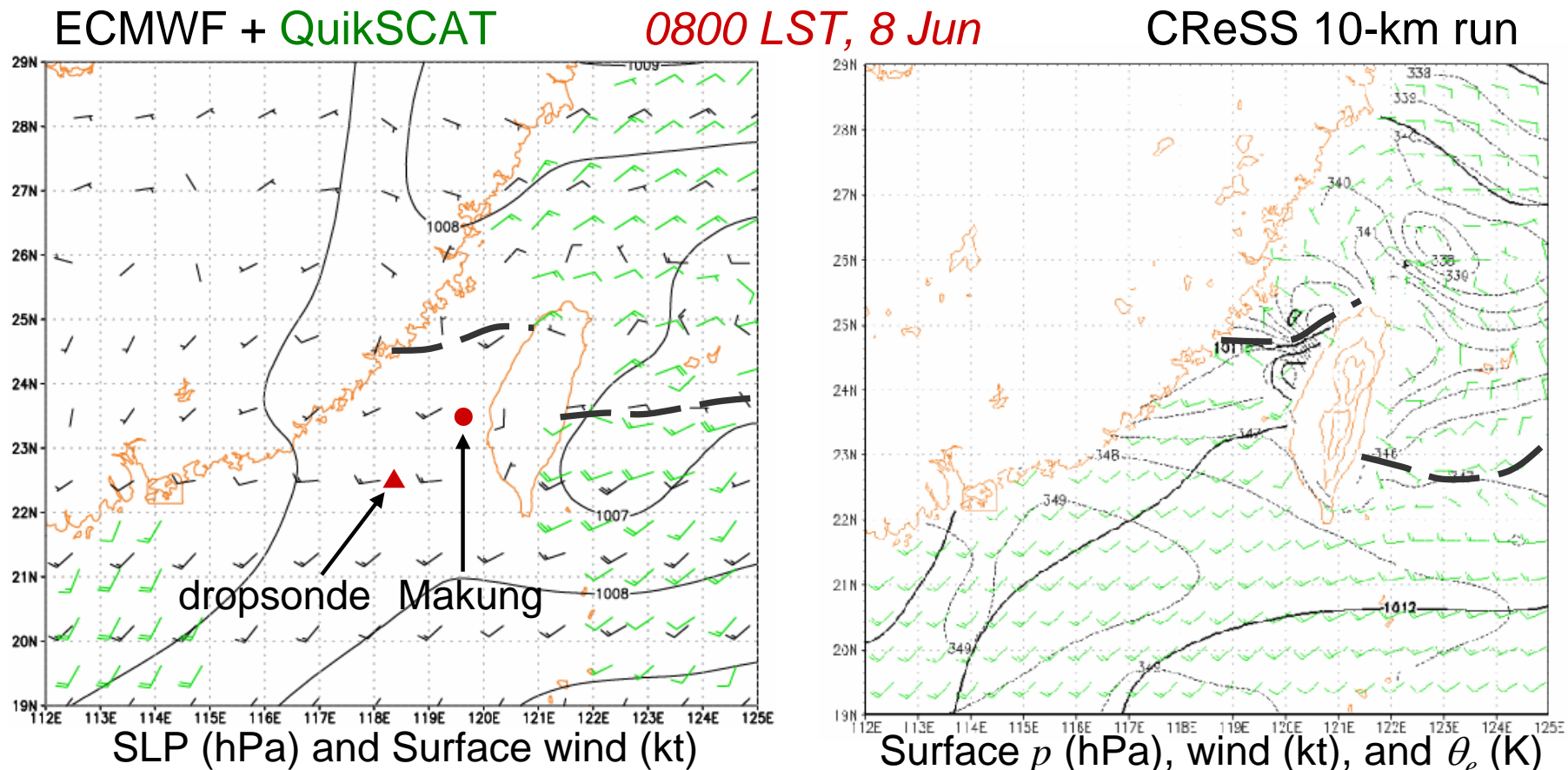
5. Simulation Result: (1)

- Surface frontal position: ECMWF + QuikSCAT vs 10-km run



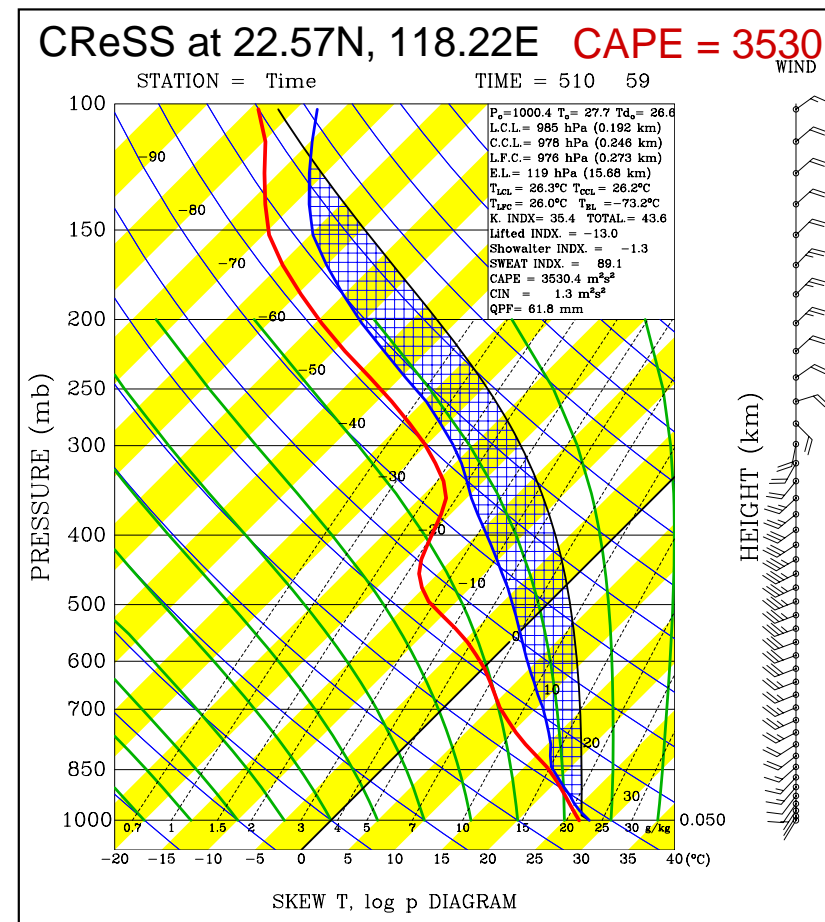
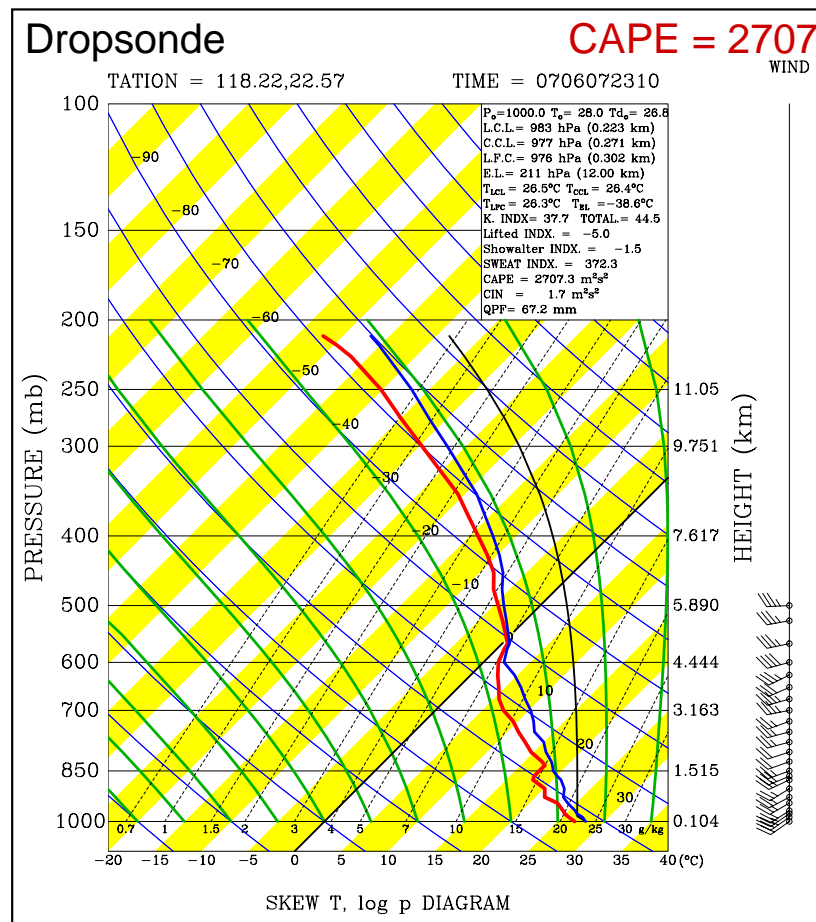
5. Simulation Result: (1)

- Surface frontal position: ECMWF + QuikSCAT vs 10-km run



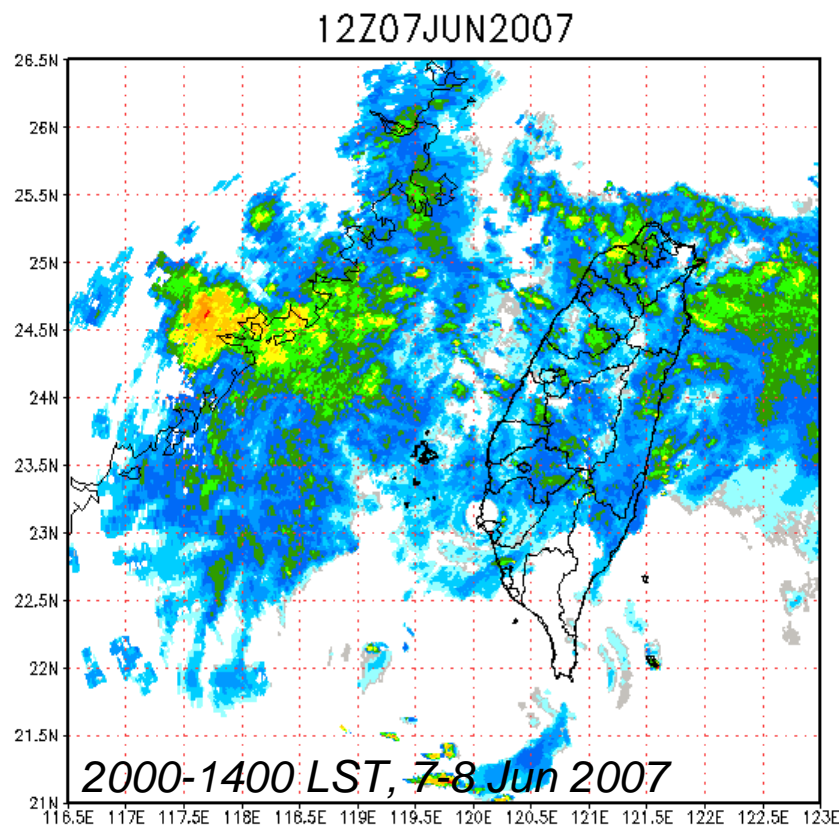
5. Simulation Result: (3)

- Sounding comparison: with dropsonde (2310 LST, 7 Jun 2007)

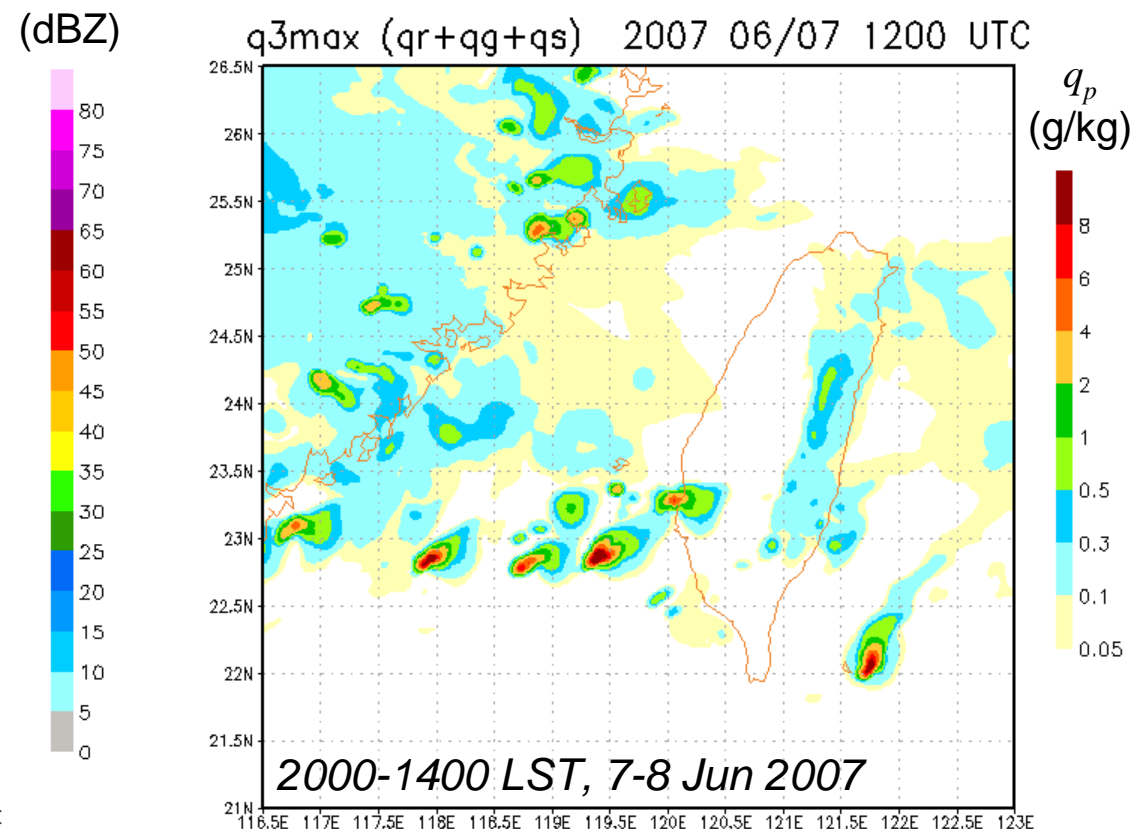


5. Simulation Result: (4)

- Convection distribution and evolution: Radar vs 2-km run

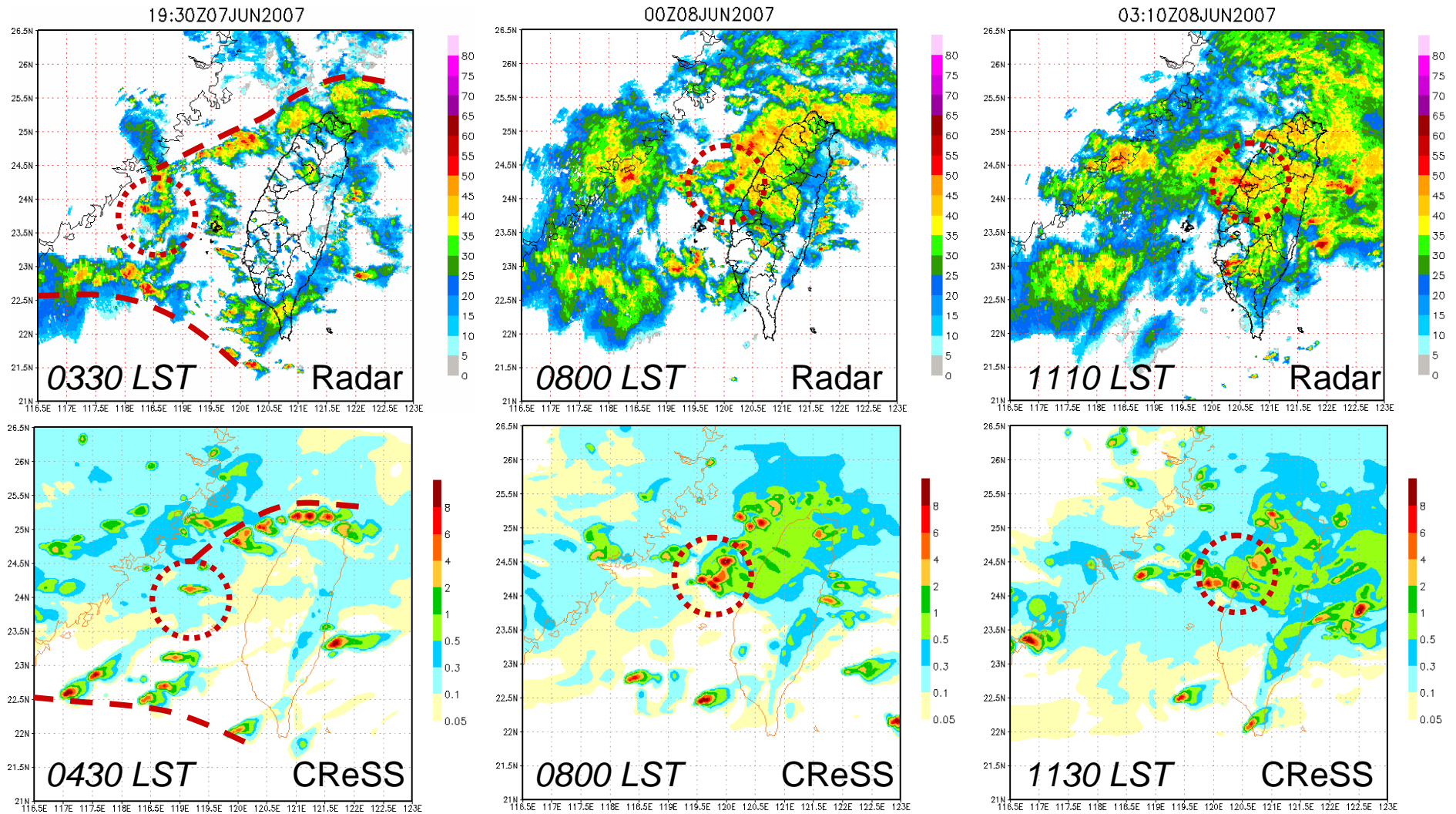


Radar Column Max Echo Composite (dBZ)



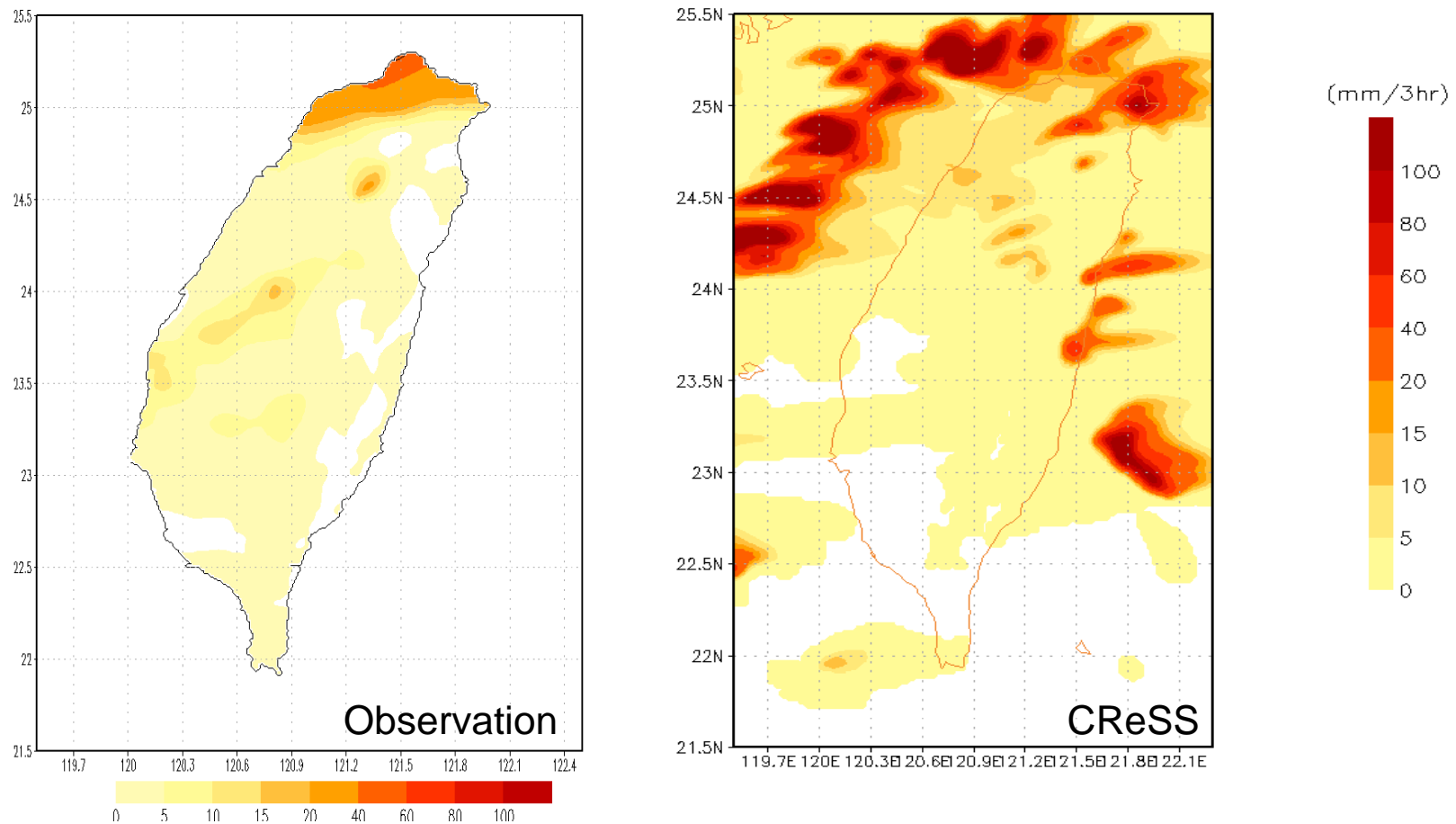
Max mixing ratio of precipitating hydrometeors

5. Simulation Result: (5)



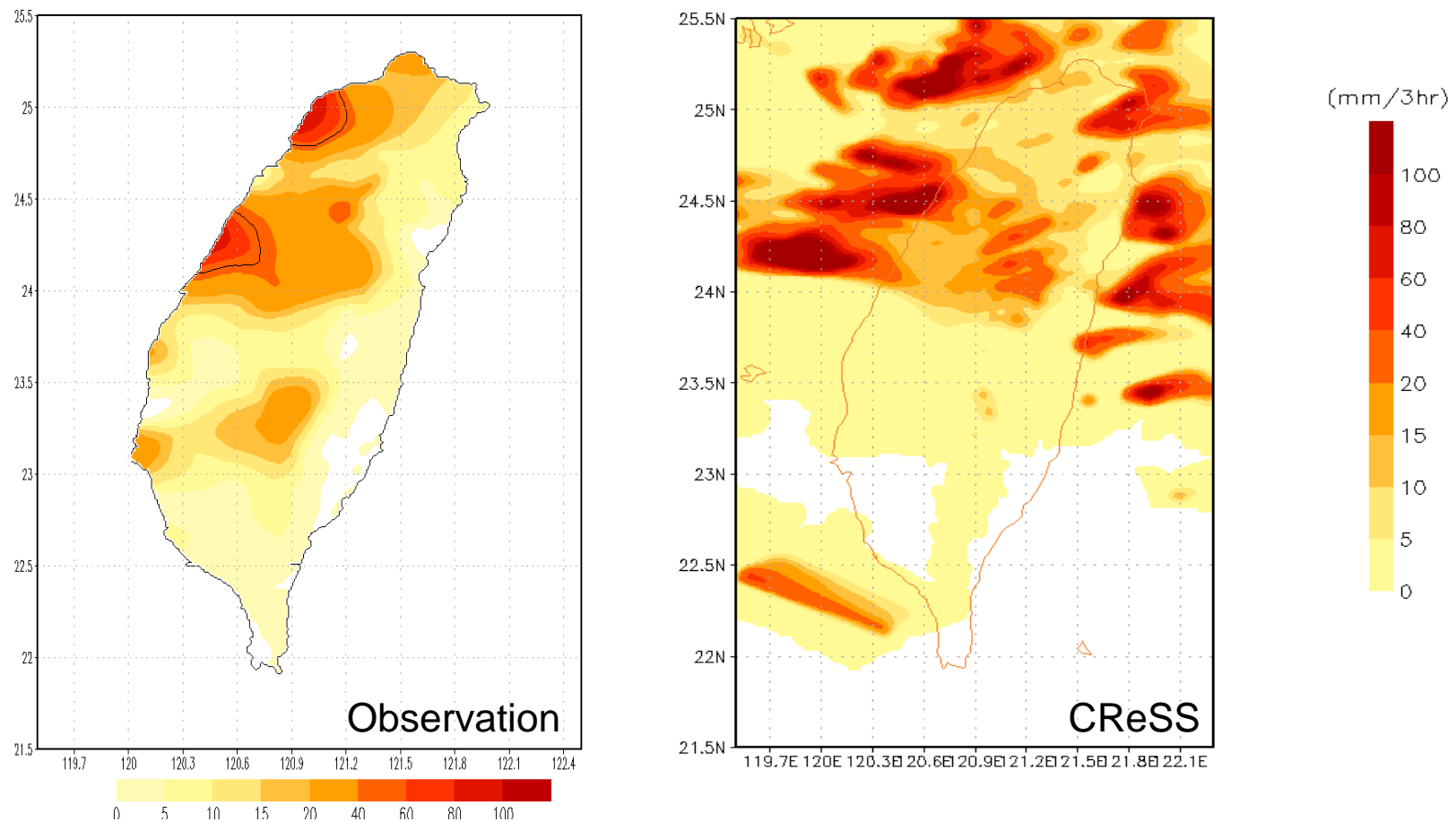
5. Simulation Result: (6)

- Rainfall accumulation every 3 h: 0500-0800 LST, 8 Jun



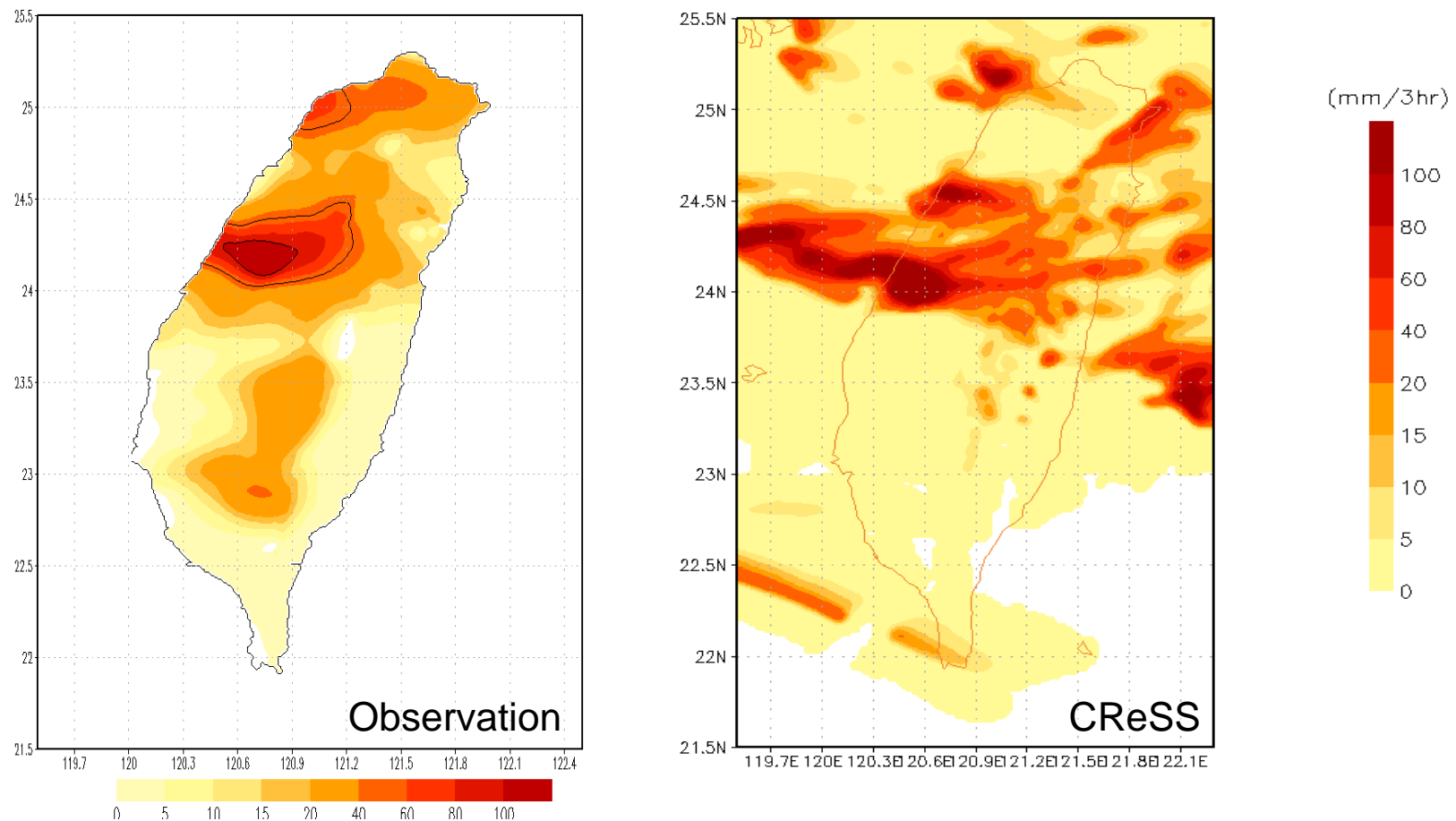
5. Simulation Result: (6)

- Rainfall accumulation every 3 h: 0800-1100 LST, 8 Jun



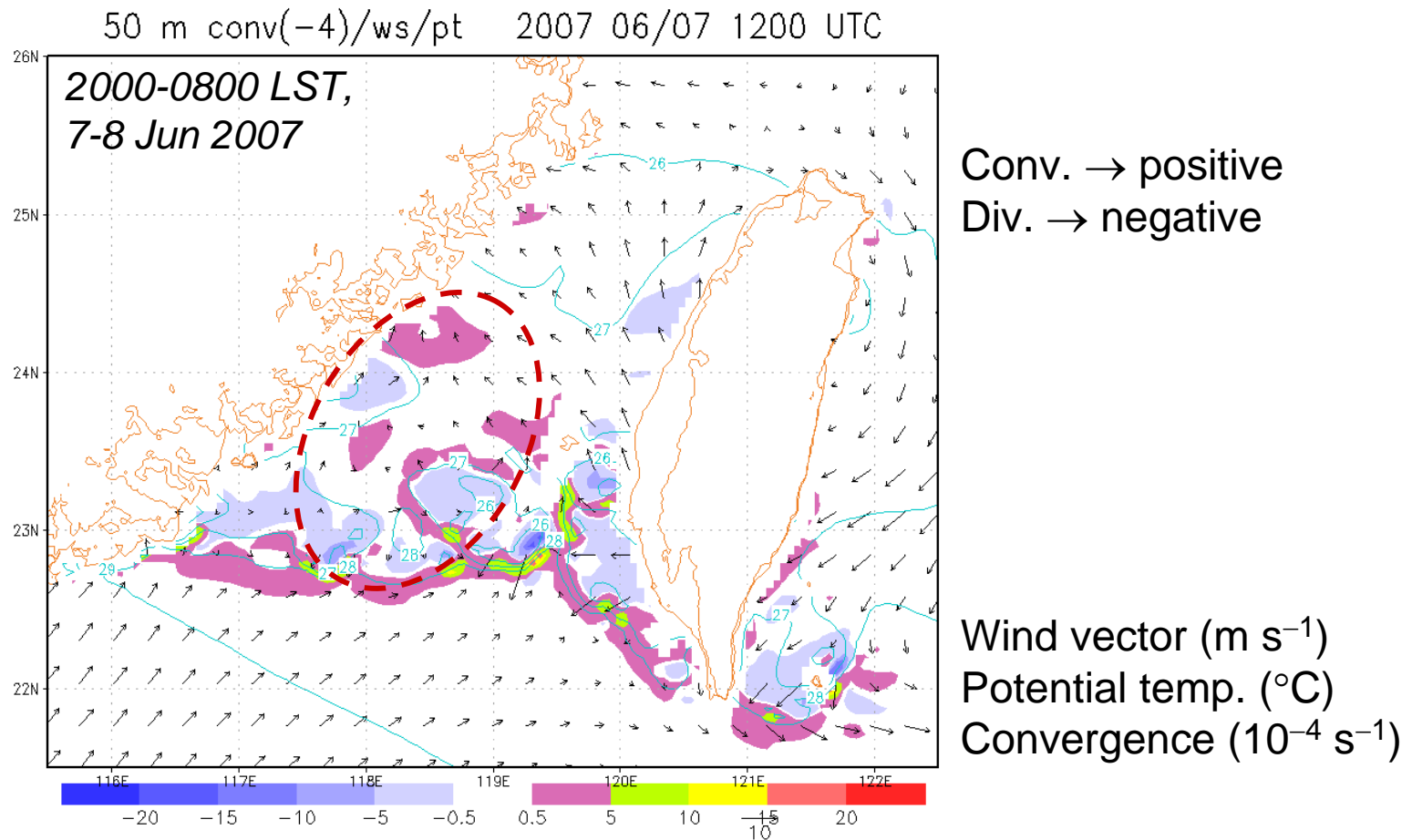
5. Simulation Result: (6)

- Rainfall accumulation every 3 h: 1100-1400 LST, 8 Jun



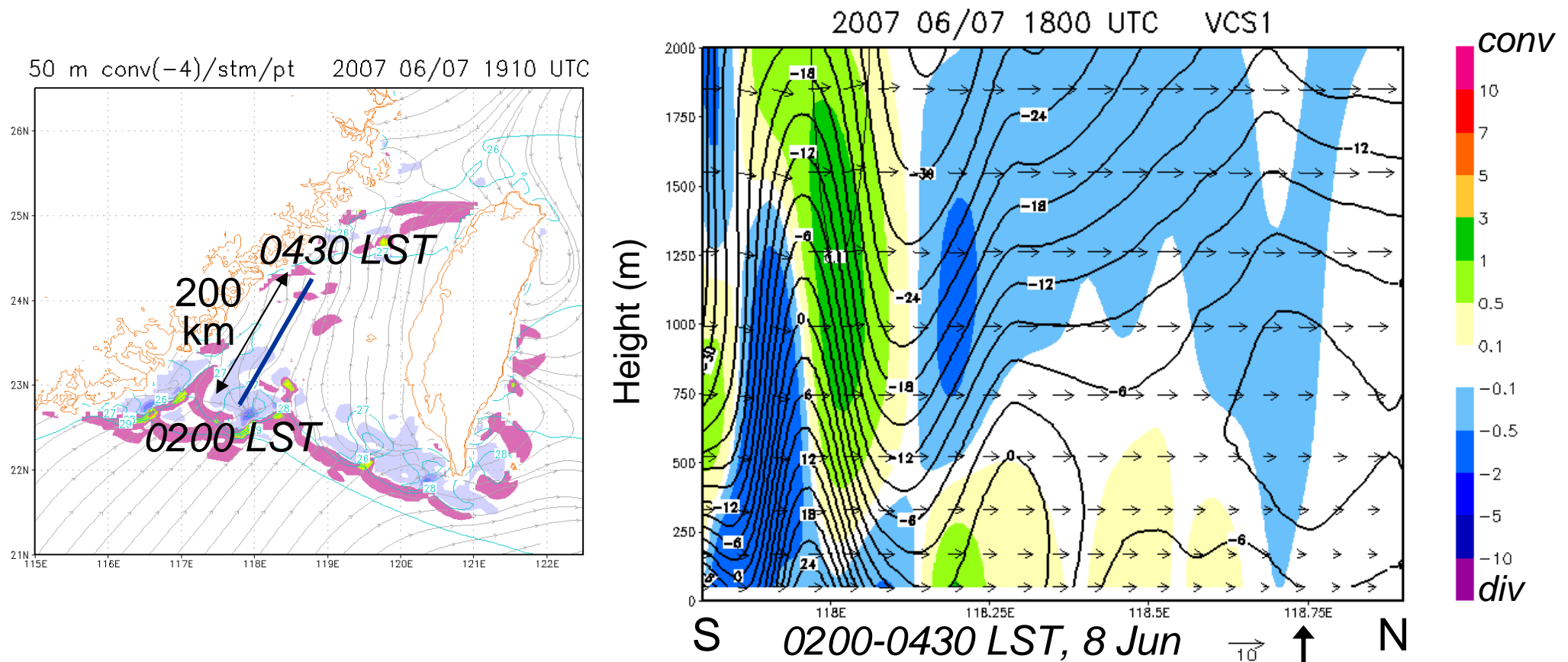
5. Simulation Result: (7)

- Wind vector, θ , and convergence at 50 m: (every 30 min)



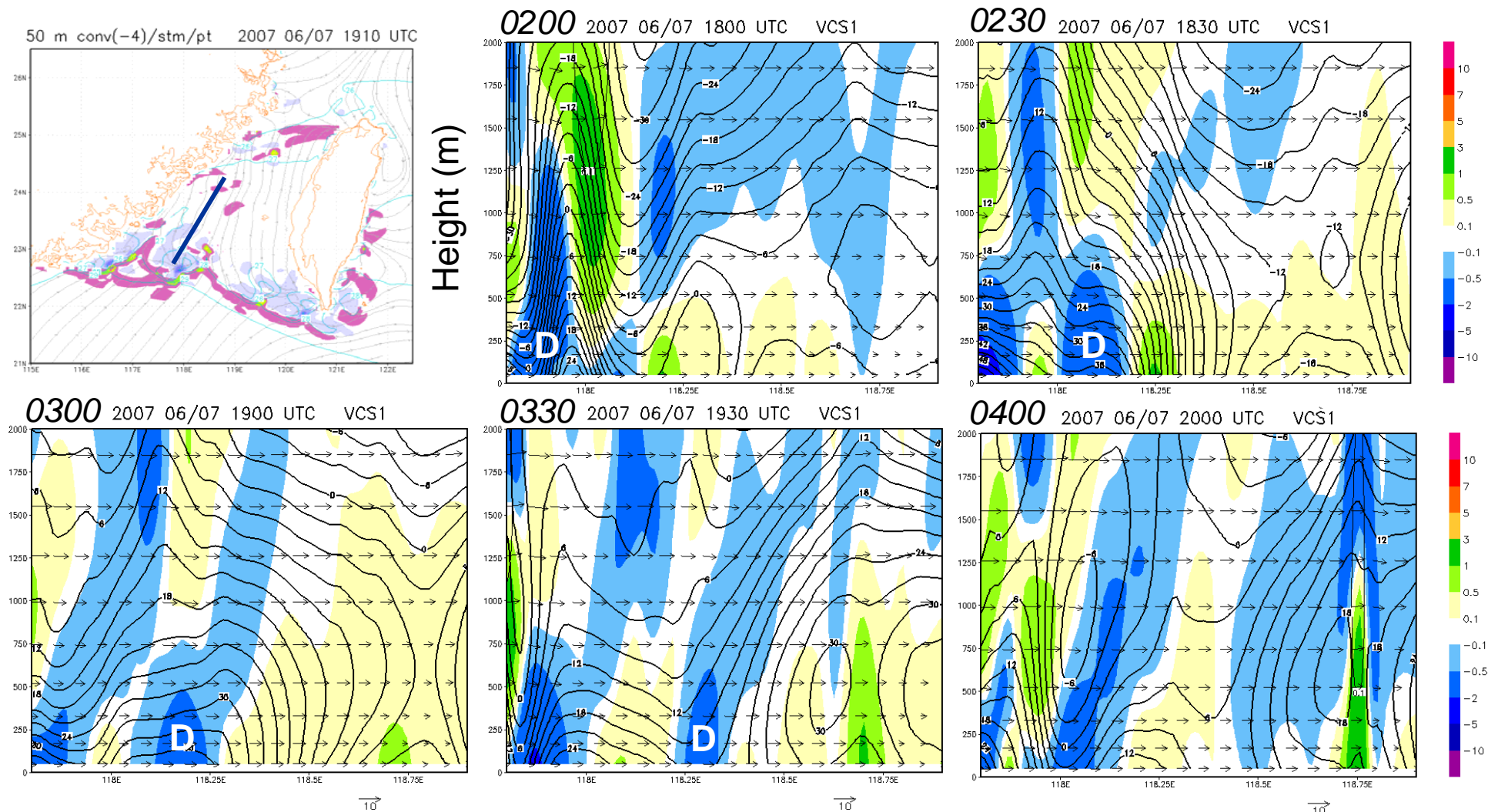
5. Simulation Result: (8)

- Vertical cross-section: Northward propagation of cold air



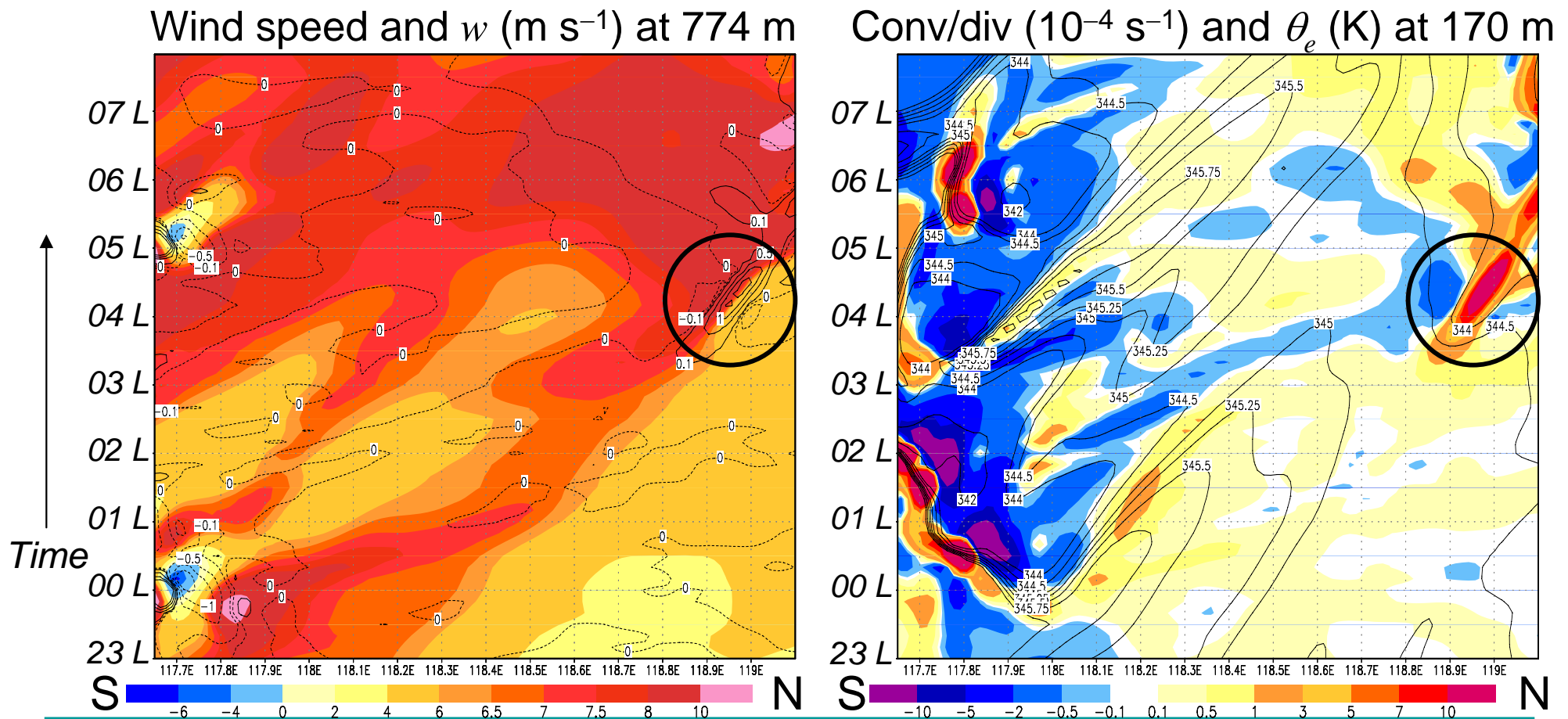
Wind vector (m s^{-1}), conv/div (10^{-4} s^{-1} , color), pressure perturbation (Pa)
Mean = 0800 LST 7 Jun to 0800 LST 8 Jun, 2007

5. Simulation Result: (9)



5. Simulation Result: (9)

- Vertical cross-section: Northward propagation of cold air



6. Discussion and Conclusion: (1)

- Favorable synoptic conditions in the present case:
 - ❑ Presence of surface Mei-yu front
 - ❑ Conditional instability with large CAPE ($2000 \text{ m}^2 \text{ s}^{-2}$), negligible CIN, and low height of LFC ($< 350 \text{ m}$)
 - ❑ LLJ at 850/700 hPa, providing vertical shear at low levels
 - ❑ Diffluent flow pattern in upper troposphere

 - Initiation and evolution of convection well reproduced in the 2-km CReSS model simulation
 - ❑ Convection along Mei-yu front (ENE-WSW) over northern strait
 - ❑ Convection from southern China (ESE-WNW) in southern strait
 - ❑ Initiation and subsequent development of convection over central strait, responsible for heavy rainfall in central Taiwan
-

6. Discussion and Conclusion: (2)

- Outflow from previous/existing convection in southern strait
 - ❑ Cold air with $D > 0$ and $p' > 0$ propagated northward
 - ❑ Increase in southerly wind speed upstream from the area of initiation, with enhanced convergence ahead
 - ❑ Lead to initiation of convection over central Taiwan Strait, and eventually heavy rainfall over central Taiwan

$$\frac{1}{\rho} \nabla p \approx \frac{1}{1.2} \frac{30}{50 \times 10^3} \approx 0.5 \times 10^{-3} \text{ m s}^{-2} \quad \Rightarrow \text{Reaching } 5 \text{ m s}^{-1} \text{ in } 3 \text{ h}$$

$$C \approx u_0 + \left(\frac{\Delta\theta}{\theta_0} g h \right)^{0.5} \approx 0 + \left(\frac{1}{100} \cdot 10 \cdot 500 \right)^{0.5} \approx 7 \text{ m s}^{-1} \quad \Rightarrow \text{Traveling } 75 \text{ km in } 3 \text{ h}$$

- In other runs starting at 12 h earlier or later:
 - ❑ No convection over southern strait, and consequently no convection reproduced over central strait at the correct time
 - ❑ Convection indeed contributed by existing cells to the south

The End

Thank you for your attention!
